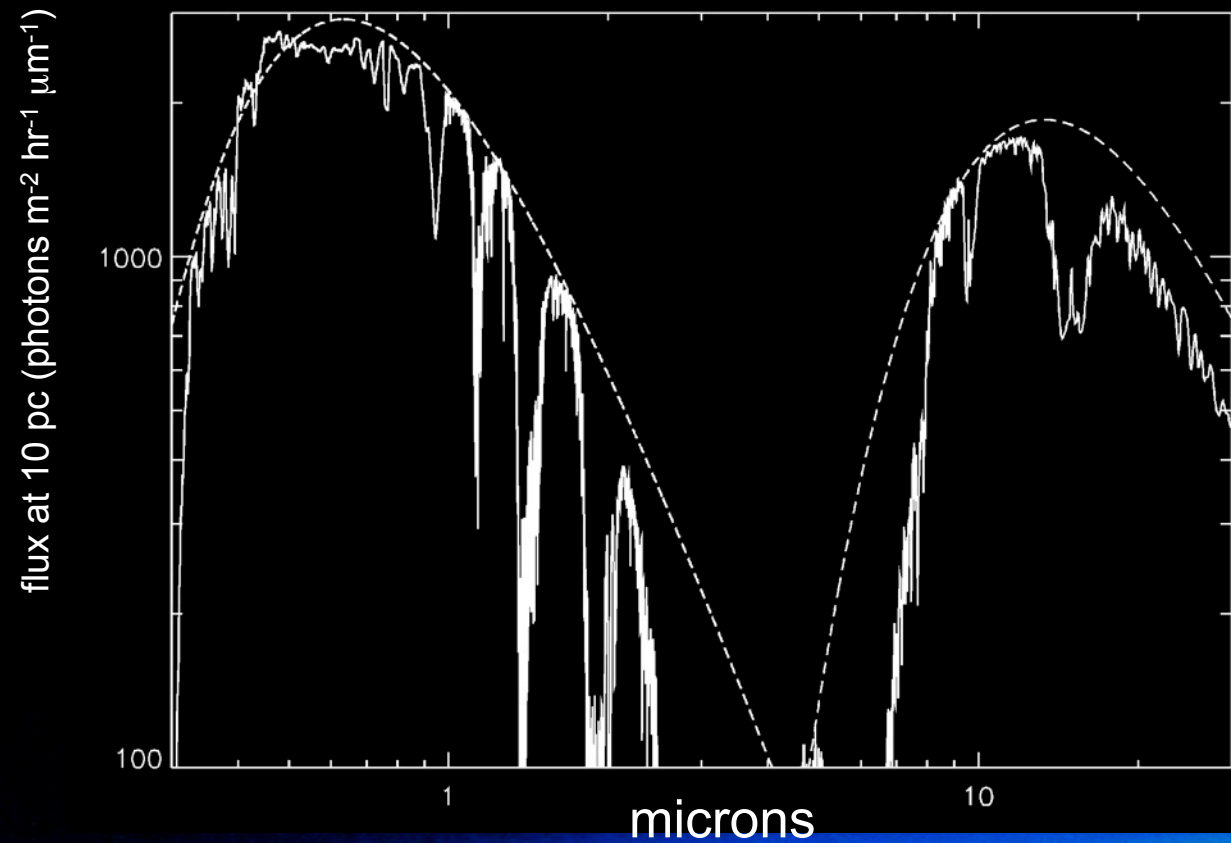


Habitable exoplanets: modeling & characterization

F. Selsis
CNRS, Bordeaux, France

R. Wordsworth, F. Forget
Laboratoire de Meteorologie
Dynamique (LMD), Paris



atmosphere: $9 \times 10^{-7} M_{\oplus}$

water: $3-5 \times 10^{-4} M_{\oplus}$

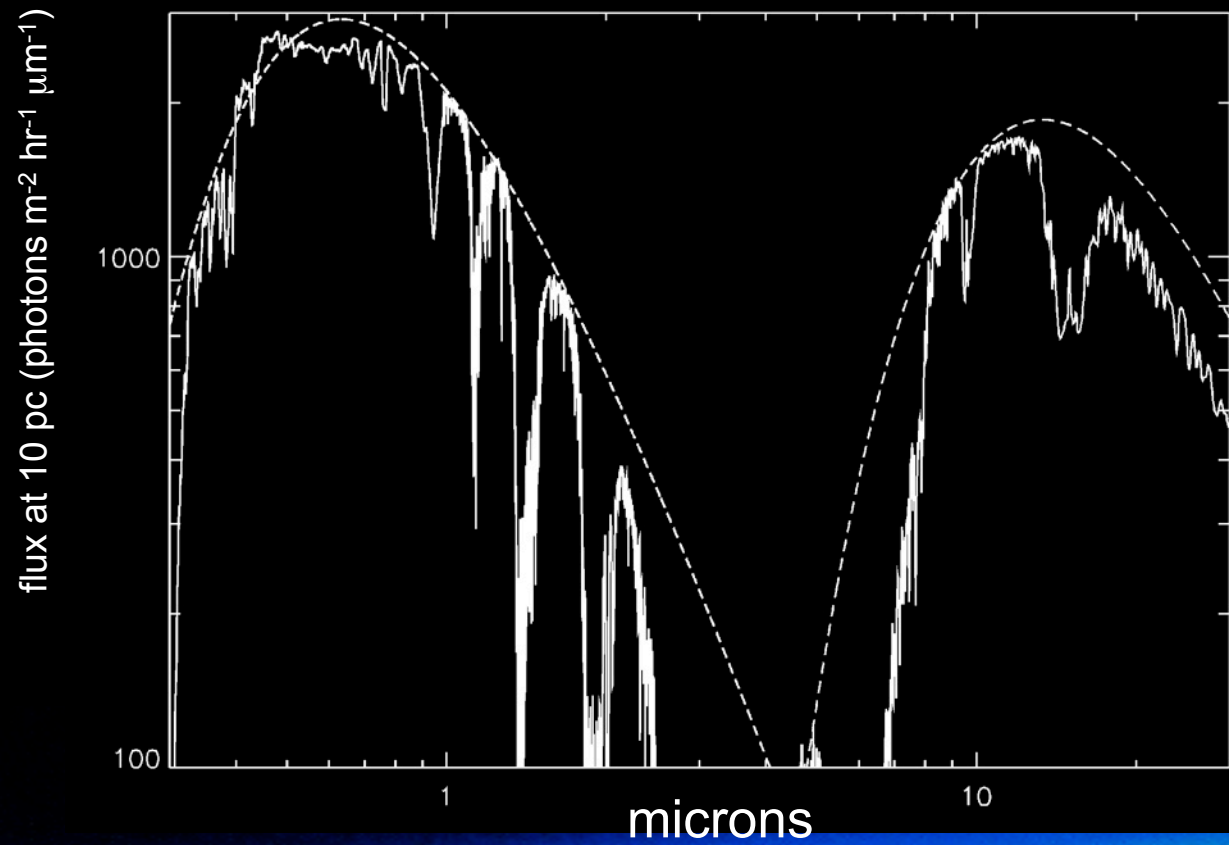
water vapor $< 1 \times 10^{-8} M_{\oplus}$

CO₂ :

$4 \times 10^{-10} M_{\oplus}$ (atmosphere)

$5 \times 10^{-5} M_{\oplus}$ (crust)

$2 \times 10^{-4} M_{\oplus}$ (mantle)



What does *habitable* mean ?

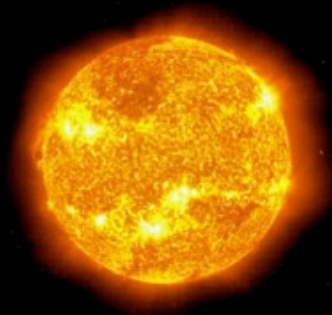
in this talk:

habitable = able to have surface liquid water

why is surface liquid water interesting ?

*** photosynthesis ***

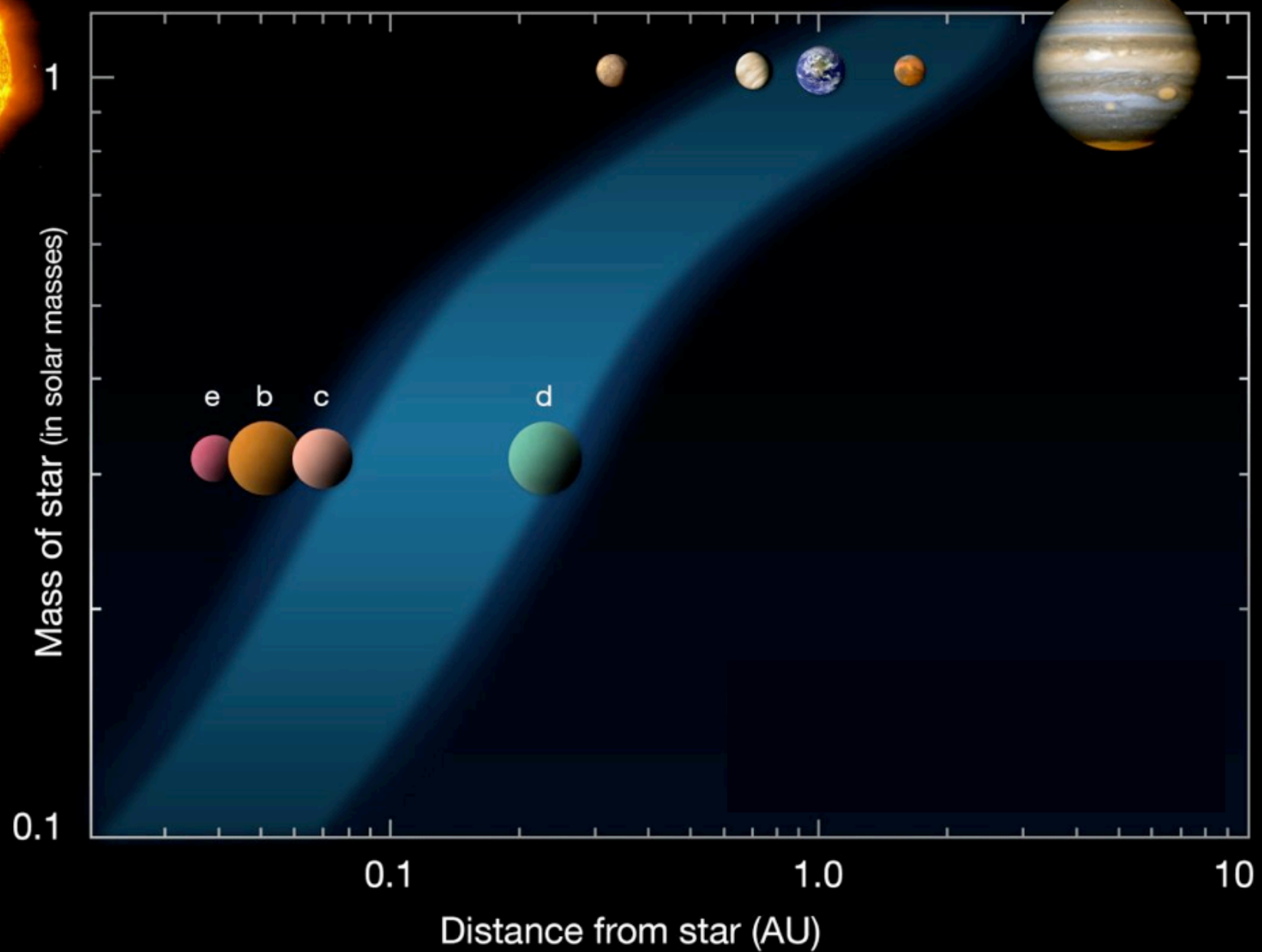




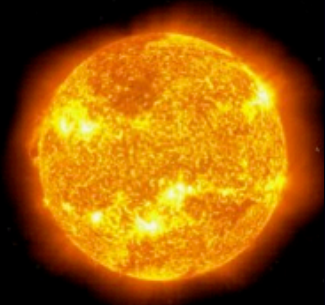
Sun



Gliese 581



Selsis et al. 2007 (HZ), Mayor et al. 2009 (Gl581 planets)

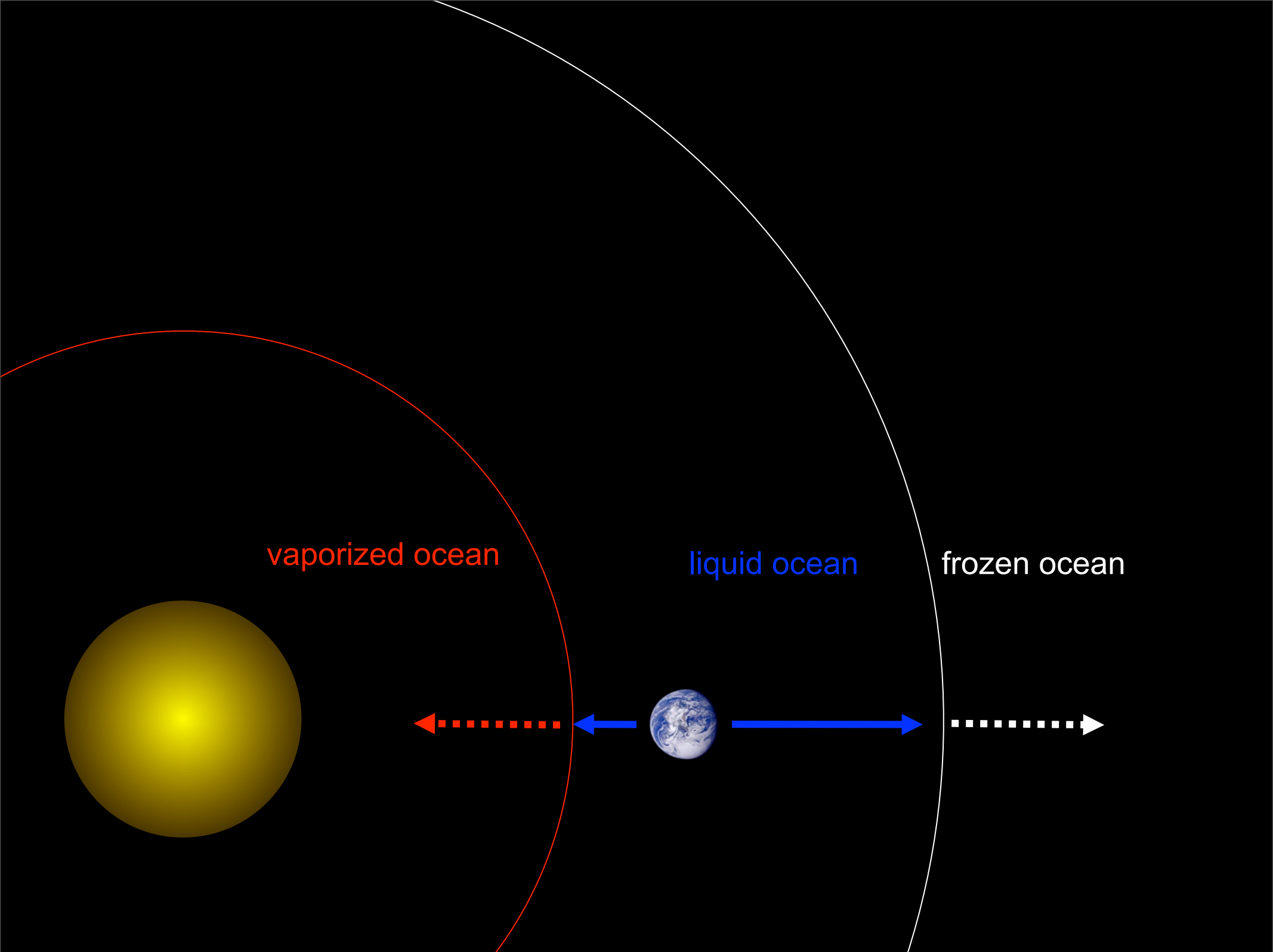


Sun



Solar HZ planets

- at least one known case !
- transiting HZ planets around G-stars will be statistically far and will offer low planet-to-star contrast ratios
- characterization requires Darwin/TPF/NWO class observatories

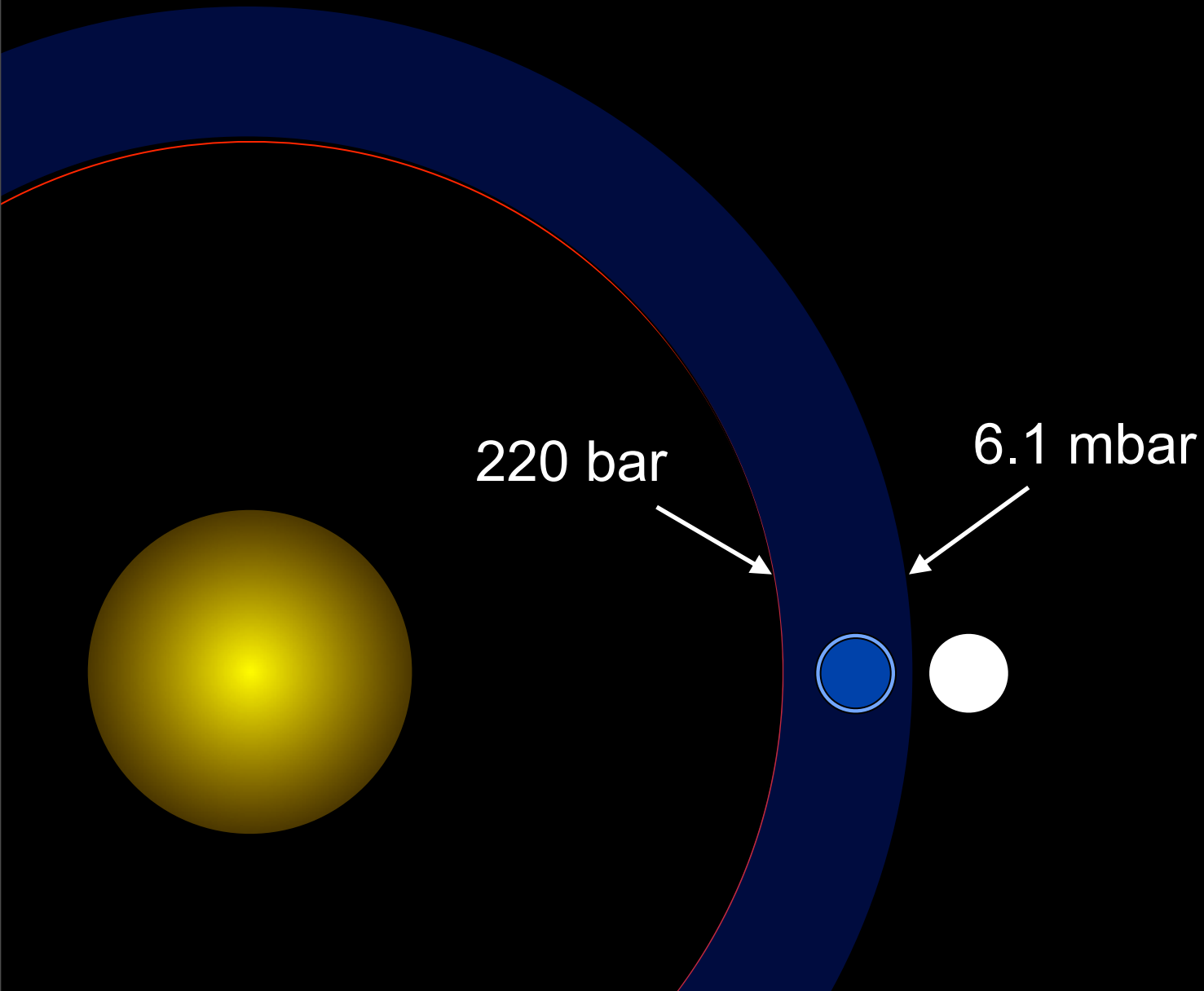


vaporized ocean

liquid ocean

frozen ocean

Inner HZ:
Water is *self-sufficient*:



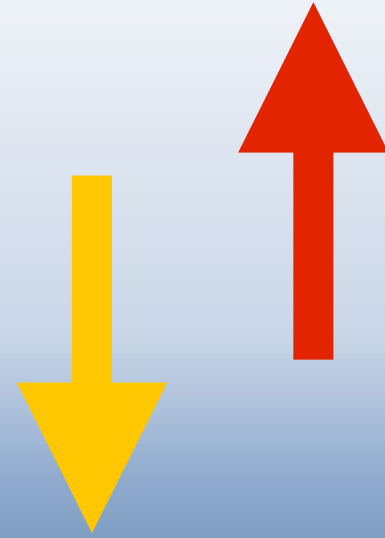
0.96 AU

T=273K (0°C)

T

log(P)

6 mbar



0.96 AU

0.93 AU

T=273K (0°C)

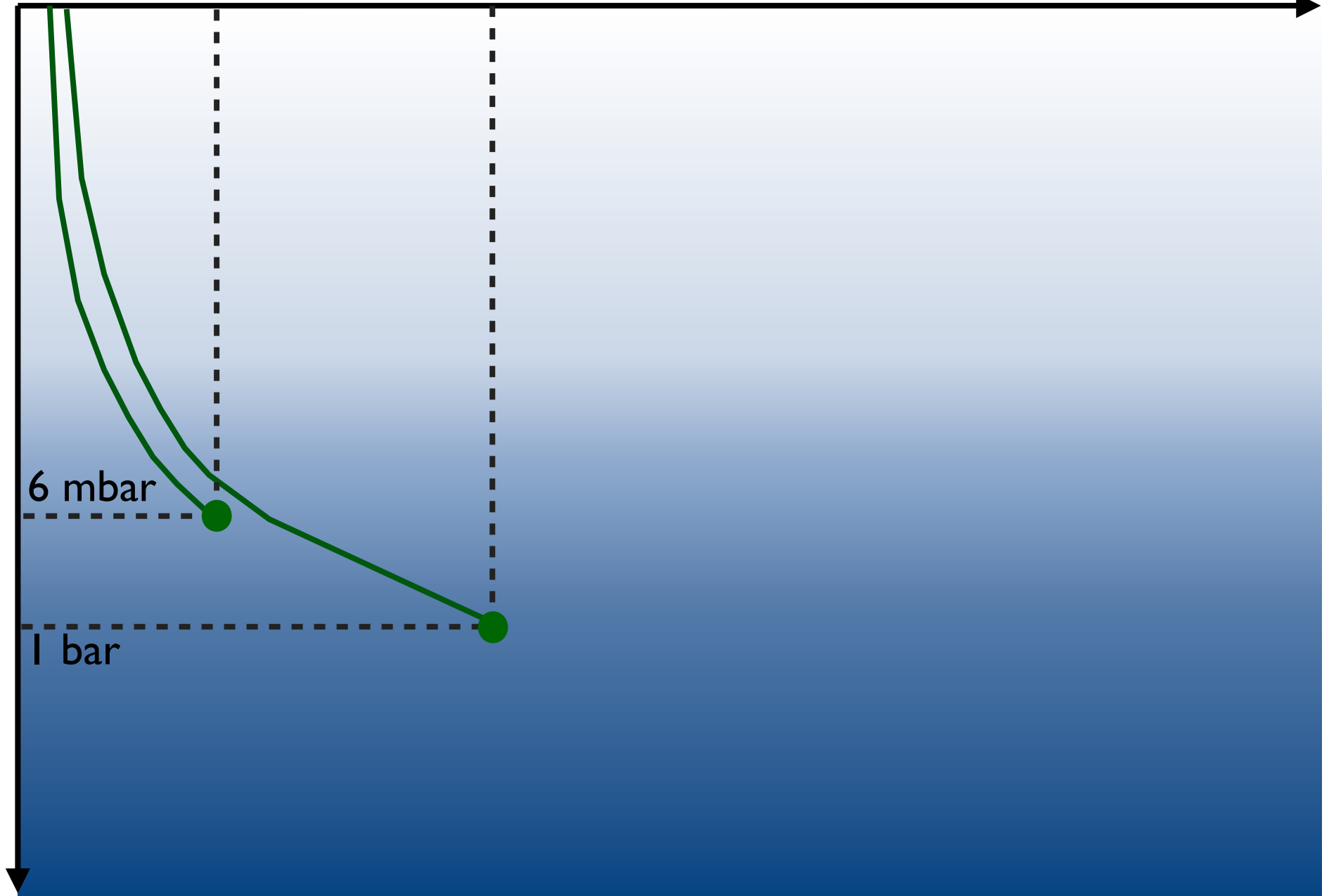
T=373K (100°C)

T

log(P)

6 mbar

1 bar



0.96 AU

0.93 AU

T=273K (0°C)

T=373K (100°C)

T

log(P)

6 mbar

1 bar



0.96 AU

0.93 AU

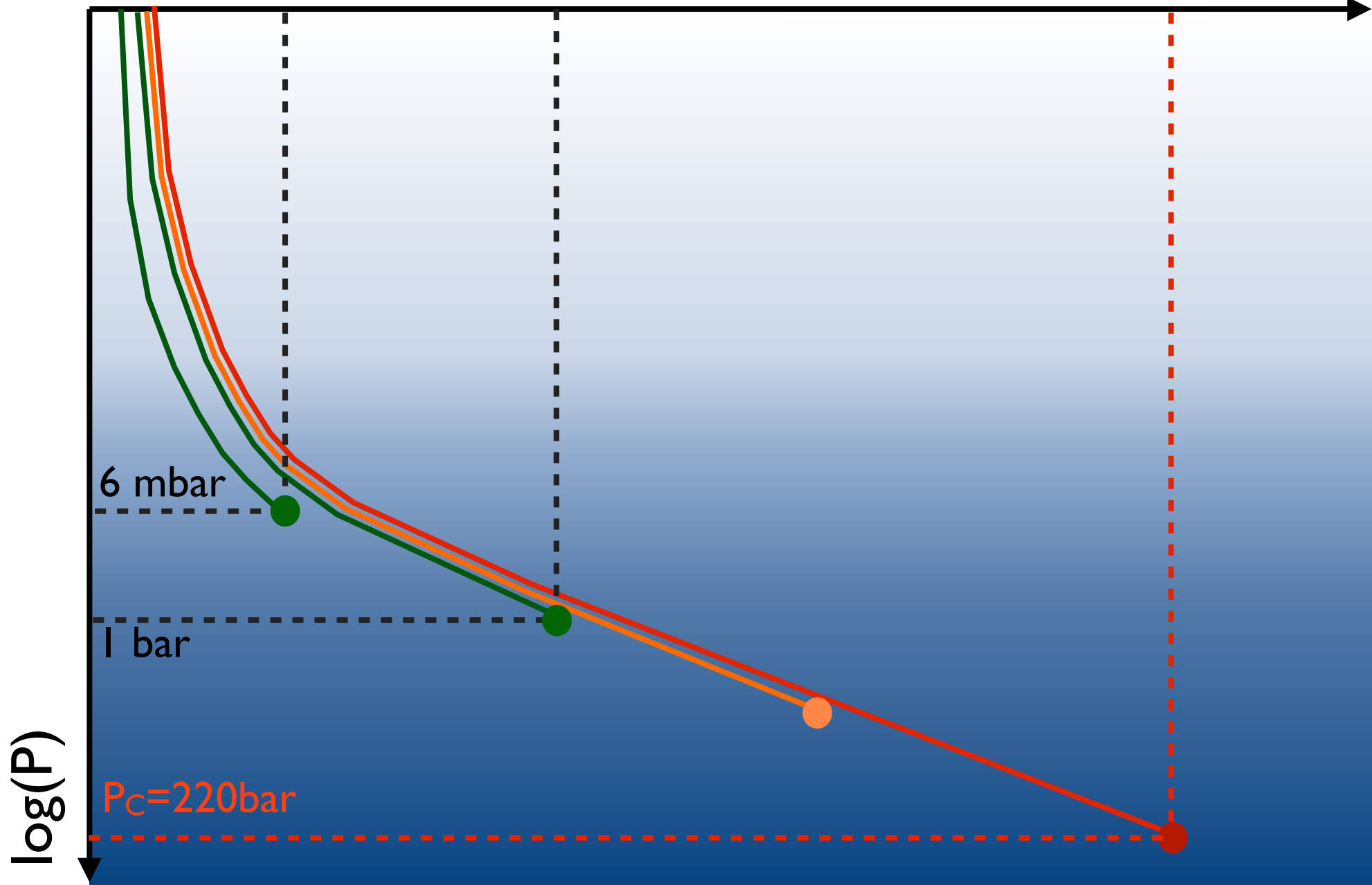
0.84 AU

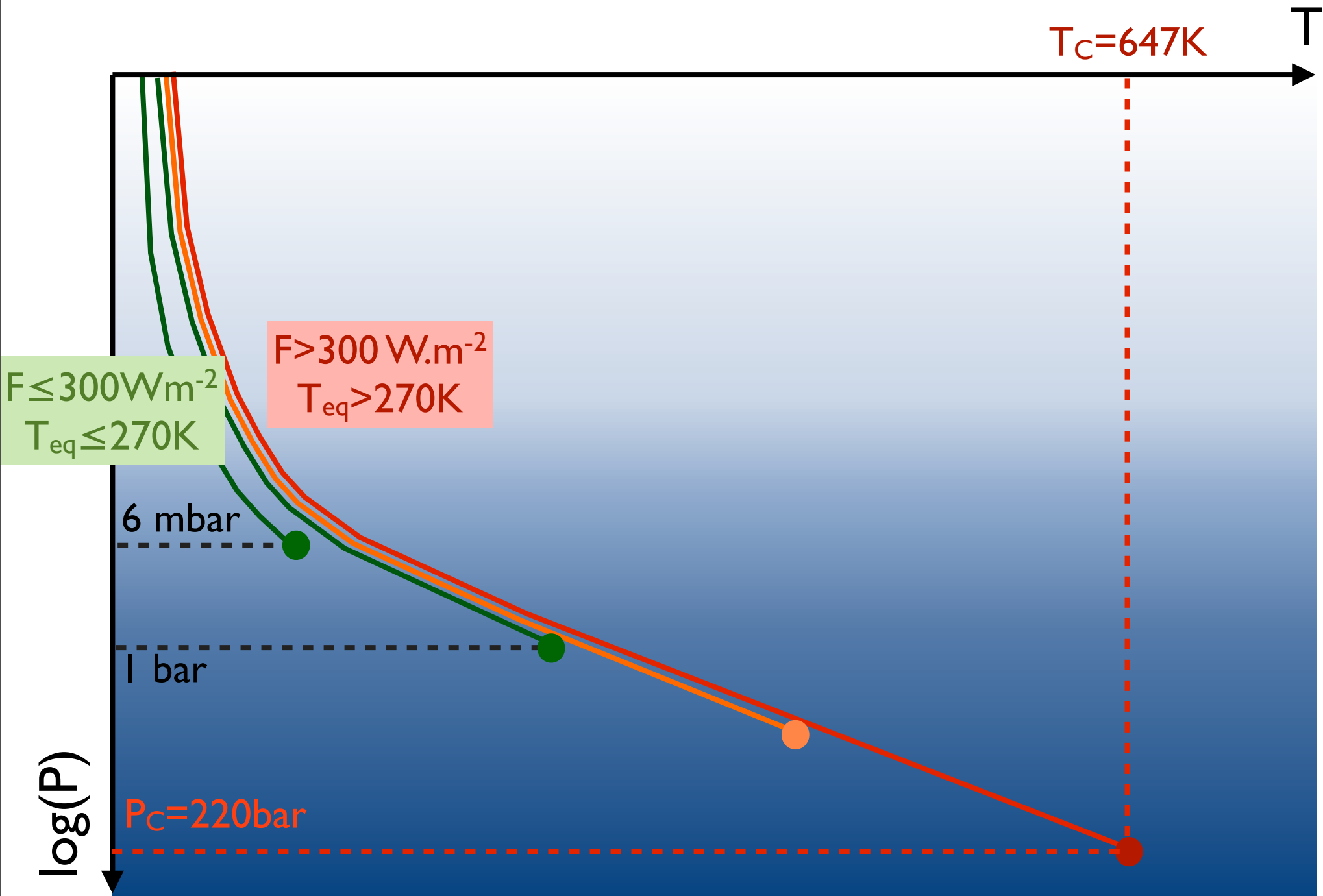
T=273K (0°C)

T=373K (100°C)

$T_C=647K$

T







$M_{\text{star}}/M_{\text{Sun}}$

Sun

(e) (b) (c)

(d)

GJ581

$T_{\text{eq}}=320\text{K}$ with $A=0.3$ (A_{Terre})
 $T_{\text{eq}}=255\text{K}$ with $A=0.75$ (A_{Venus})

$T_{\text{eq}}=255\text{K}$ with $A=0.3$

0.1

0.1

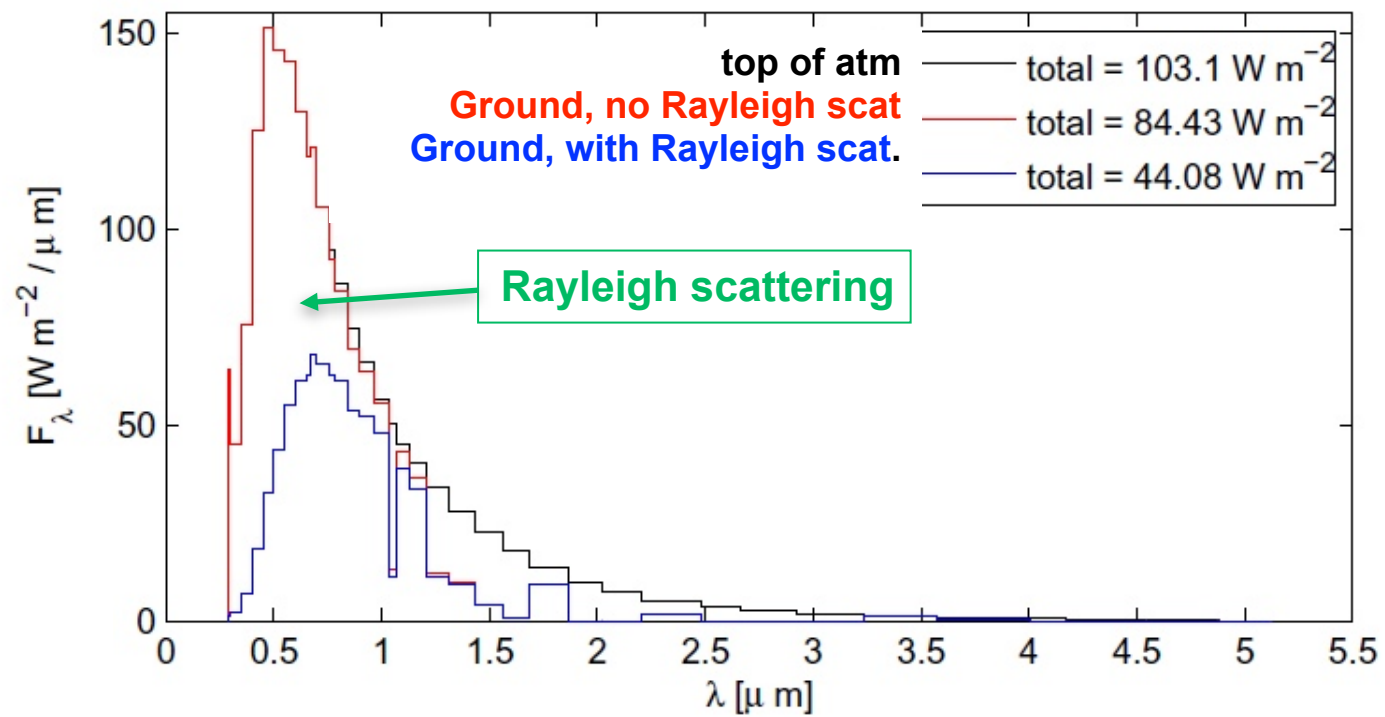
1.0

10

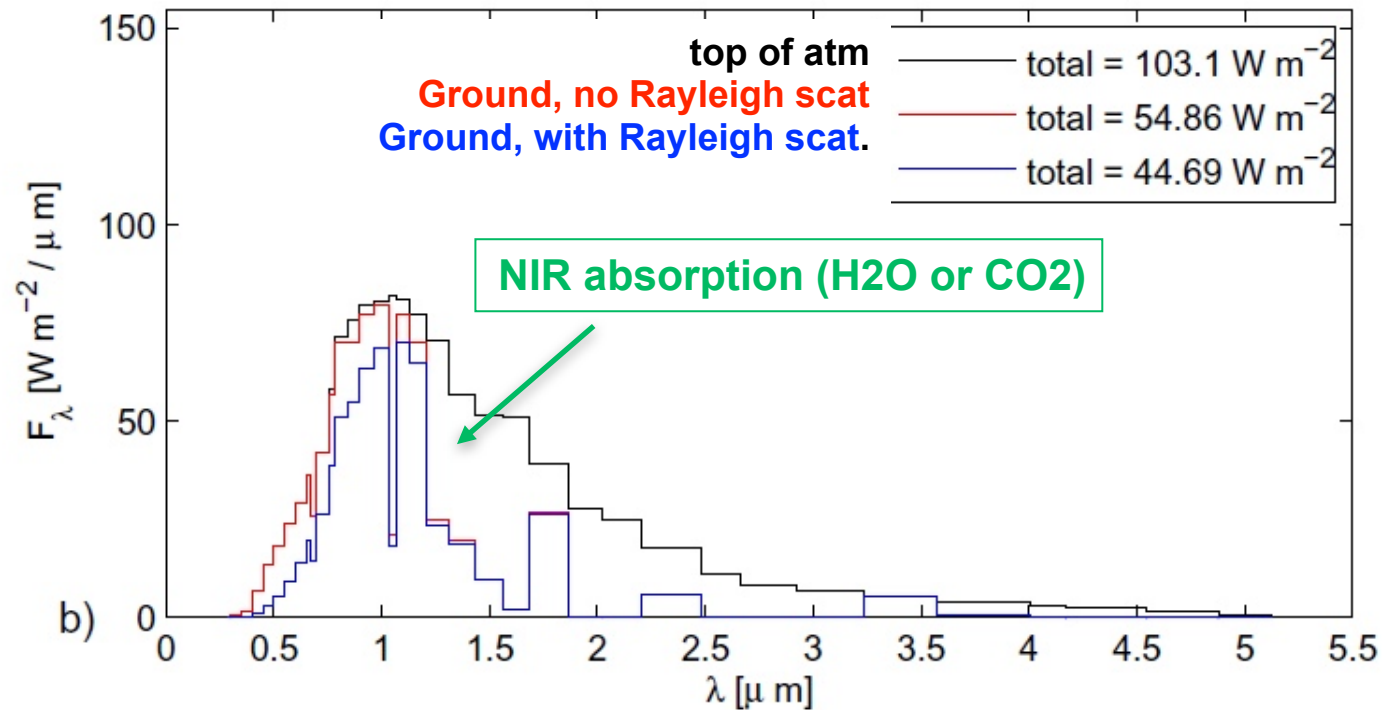
orbital distance (AU)

Vitesse Radiale GJ581: Udry et al., 2007; Mayor et al., 2009 - HZ: Selsis et al., 2007

G star



CO₂ atmosphere



M star



$M_{\text{star}}/M_{\text{Sun}}$

Sun

(e) (b) (c)

(d)

GJ581

$T_{\text{eq}}=320\text{K}$ with $A=0.3$ (A_{Terre})

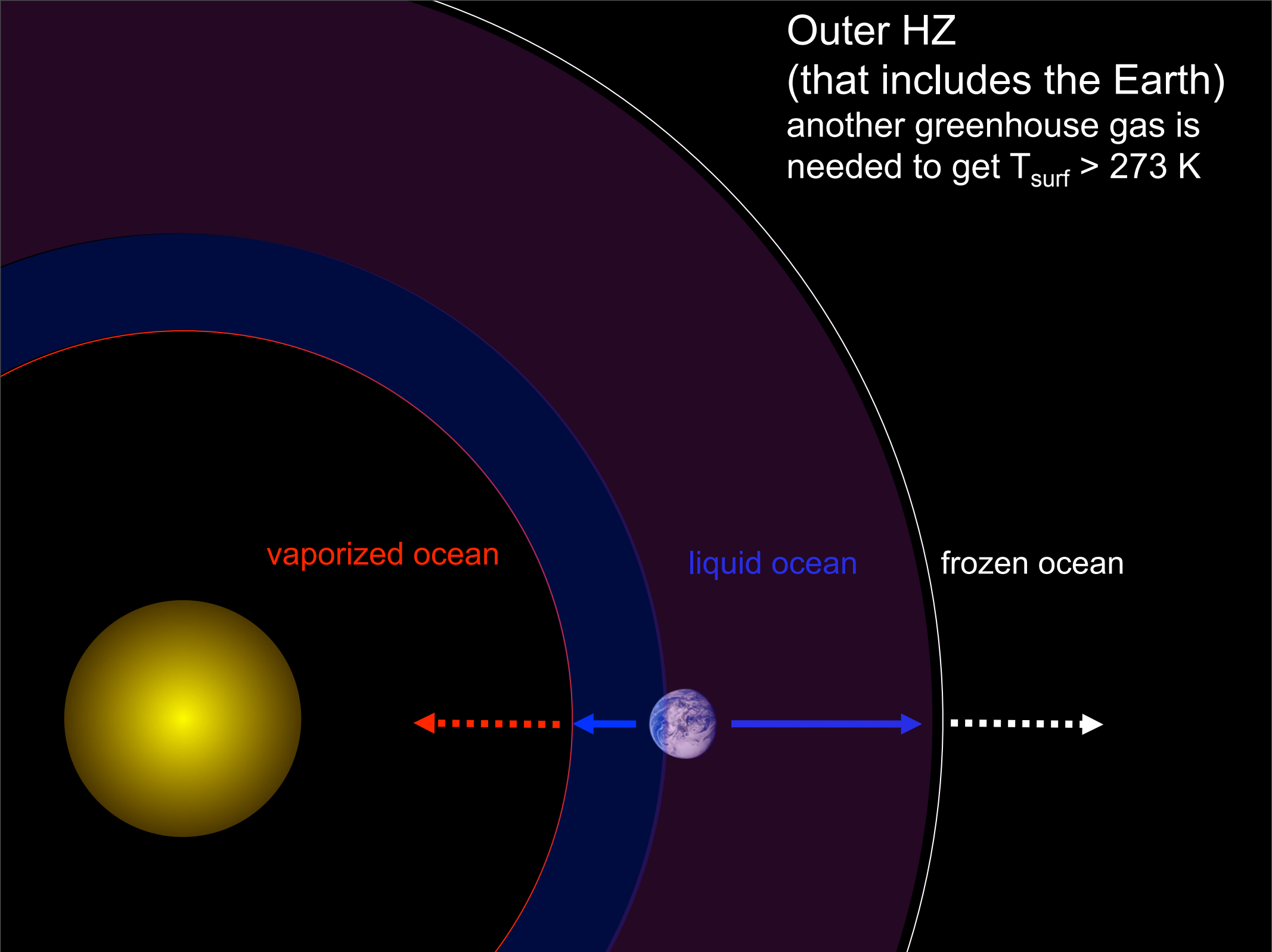
~~$T_{\text{eq}}=255\text{K}$ with $A=0.75$ (A_{Venus})~~

$T_{\text{eq}}=255\text{K}$ with $A=0.3$

orbital distance (AU)

Vitesse Radiale GJ581: Udry et al., 2007; Mayor et al., 2009 - HZ: Selsis et al., 2007

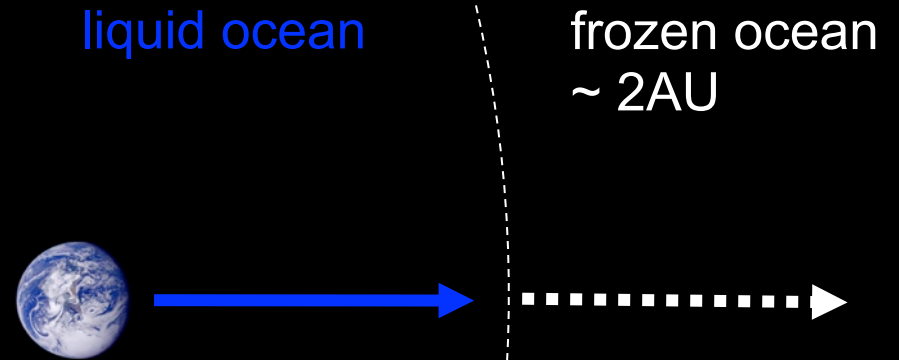
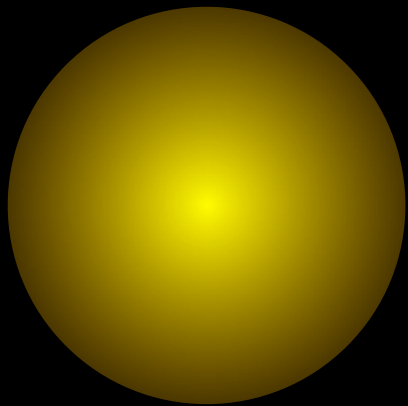
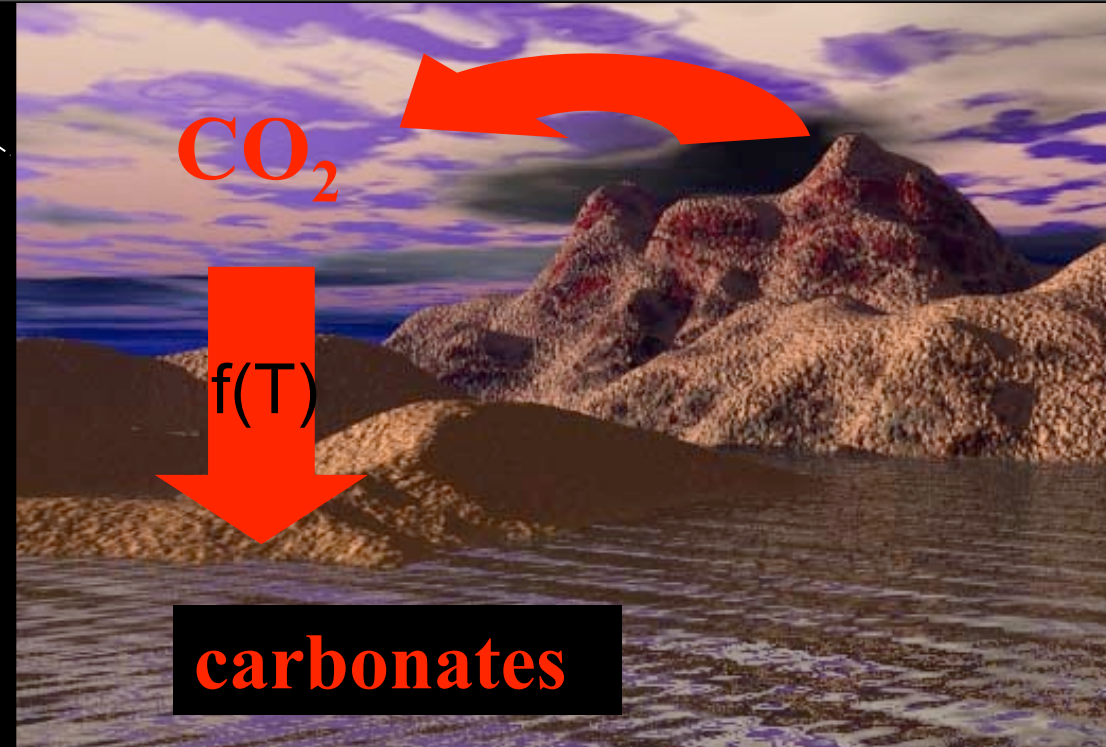
Outer HZ
(that includes the Earth)
another greenhouse gas is
needed to get $T_{\text{surf}} > 273 \text{ K}$



Outer boundary of the HZ

Walker et al., 1981 (carbonate-silicate cycle)

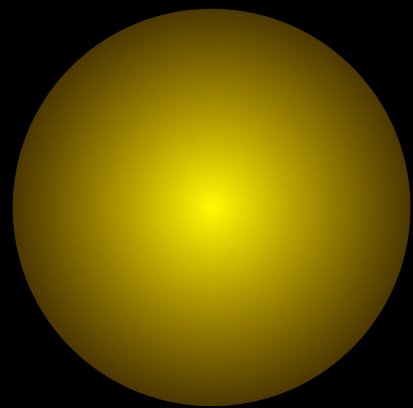
Kasting et al., 1993



Outer boundary of the HZ

Walker et al., 1981 (carbonate-
silicate cycle)

Kasting et al., 1993



liquid ocean

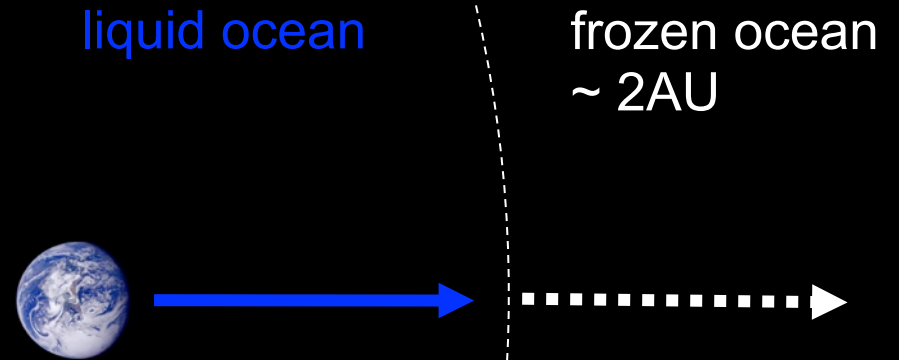
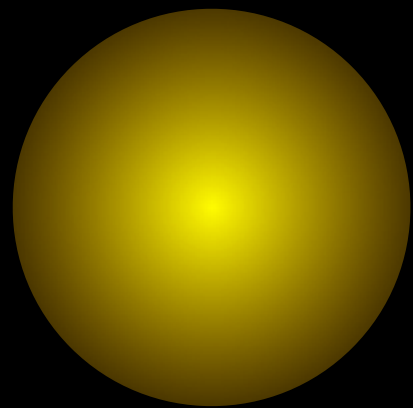
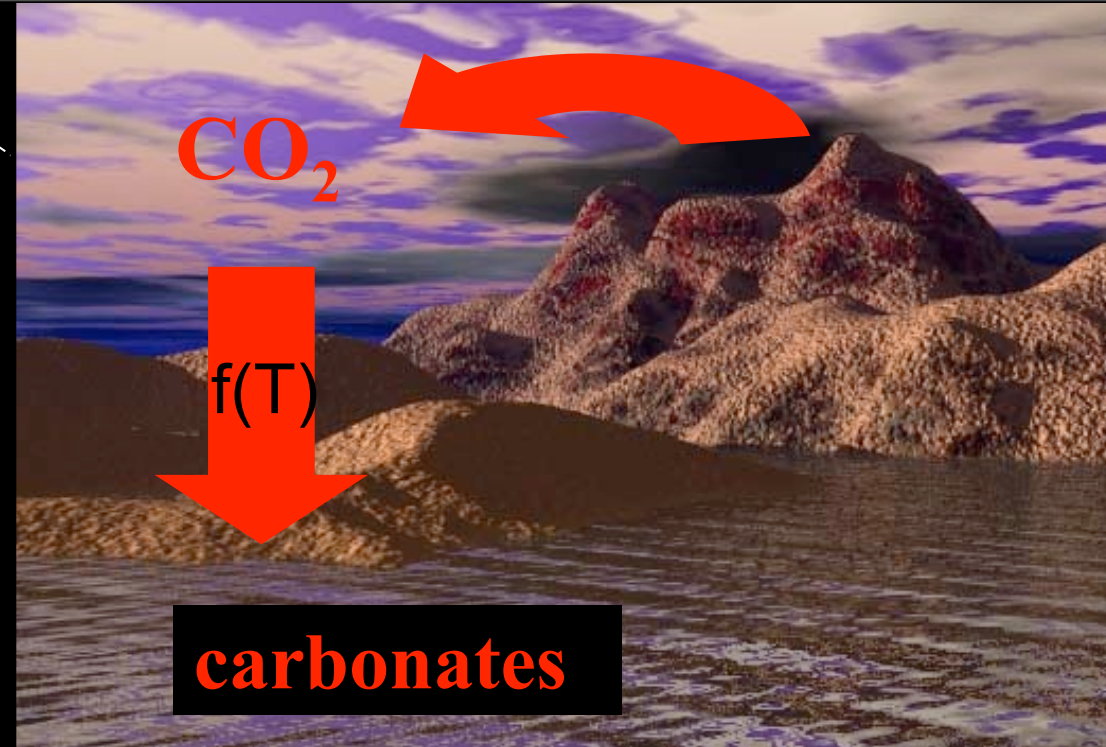
frozen ocean
~ 2AU

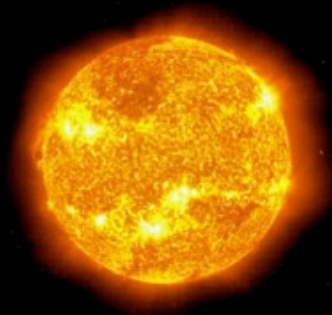


Outer boundary of the HZ

Walker et al., 1981 (carbonate-silicate cycle)

Kasting et al., 1993

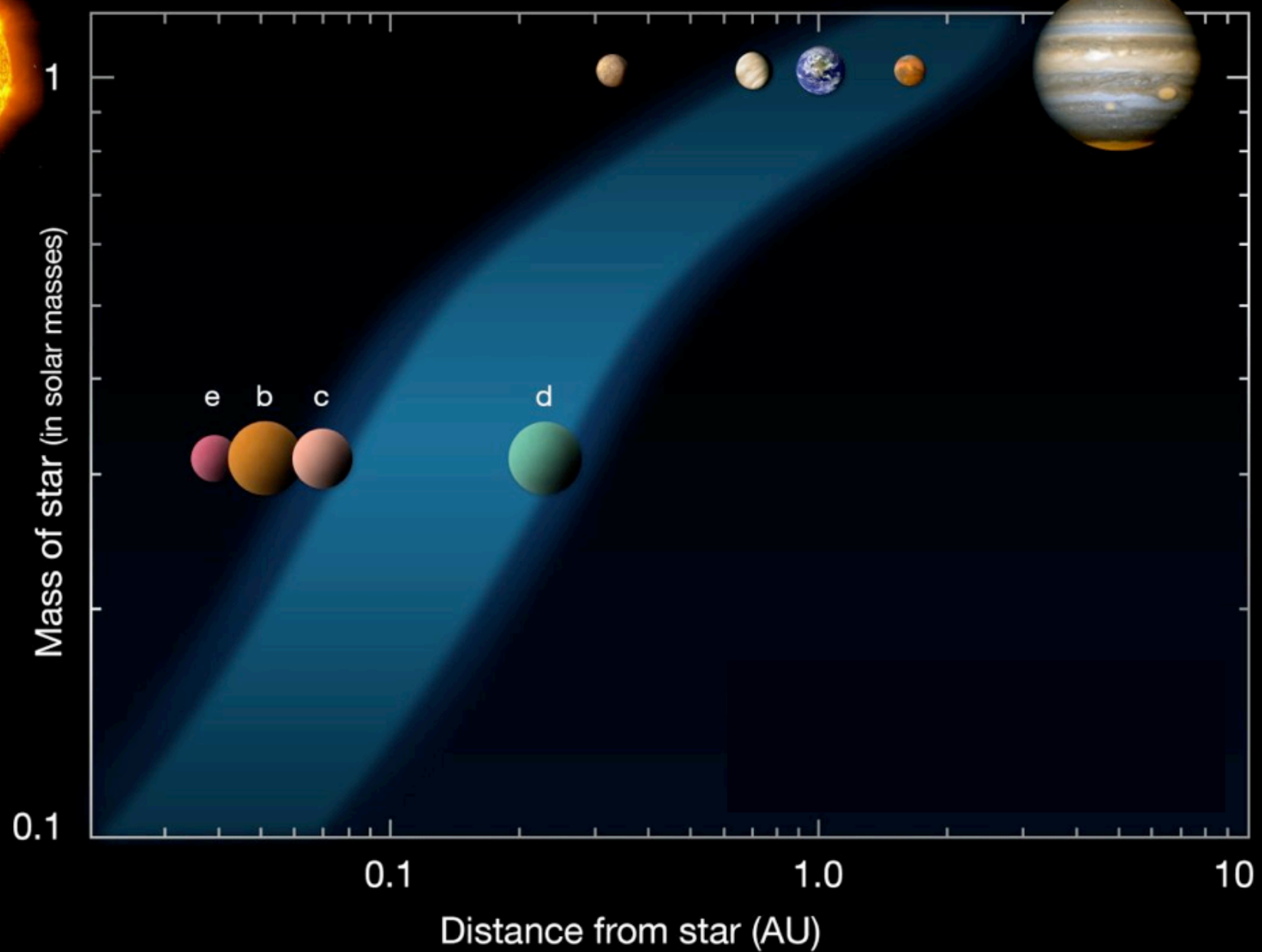




Sun



Gliese 581



Selsis et al. 2007 (HZ), Mayor et al. 2009 (Gl581 planets)

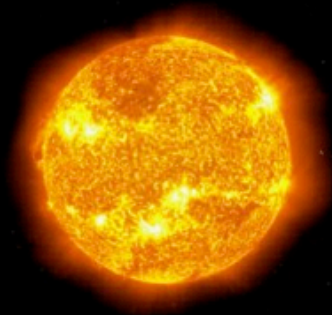
M-stars HZ planets

- most stars are M stars
- HZ planets are detectable by RV
- if transiting, the planet-to-star contrast ratio can make eclipse spectroscopy (marginally) possible



Gliese 581

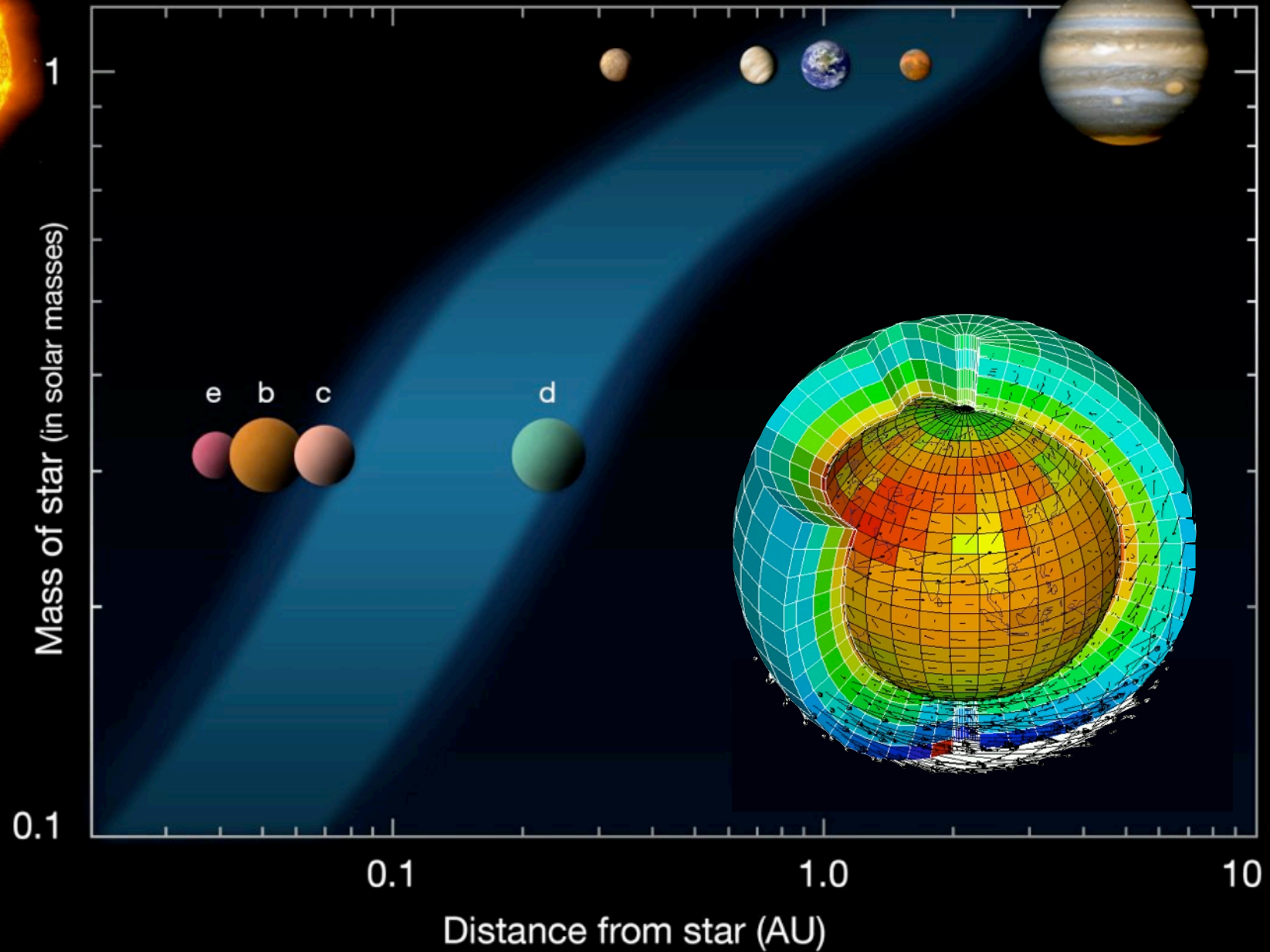
- tidally evolved (affects rotation rate and obliquity)
- different spectrum (cannot scale modeling done with solar irradiation)
- M stars remain active for a long time: strong X, EUV, stellar winds and associated atmospheric erosion



Sun



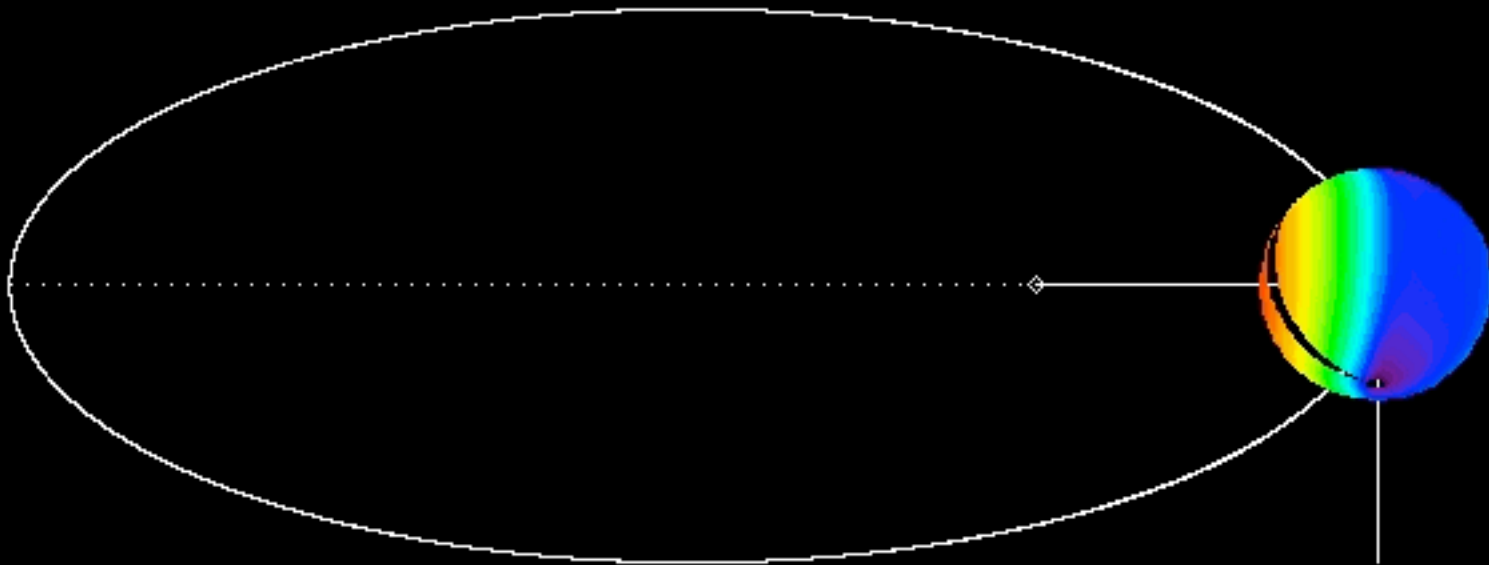
Gliese 581



Selsis et al. 2007 (HZ), Mayor et al. 2009 (Gl581 planets)

Eccentric planets

- mean insolation = $F(a)(1-e^2)^{-1/2}$
- variation of insolation
- rotation (equilibrium or resonance)
- tidal dissipation



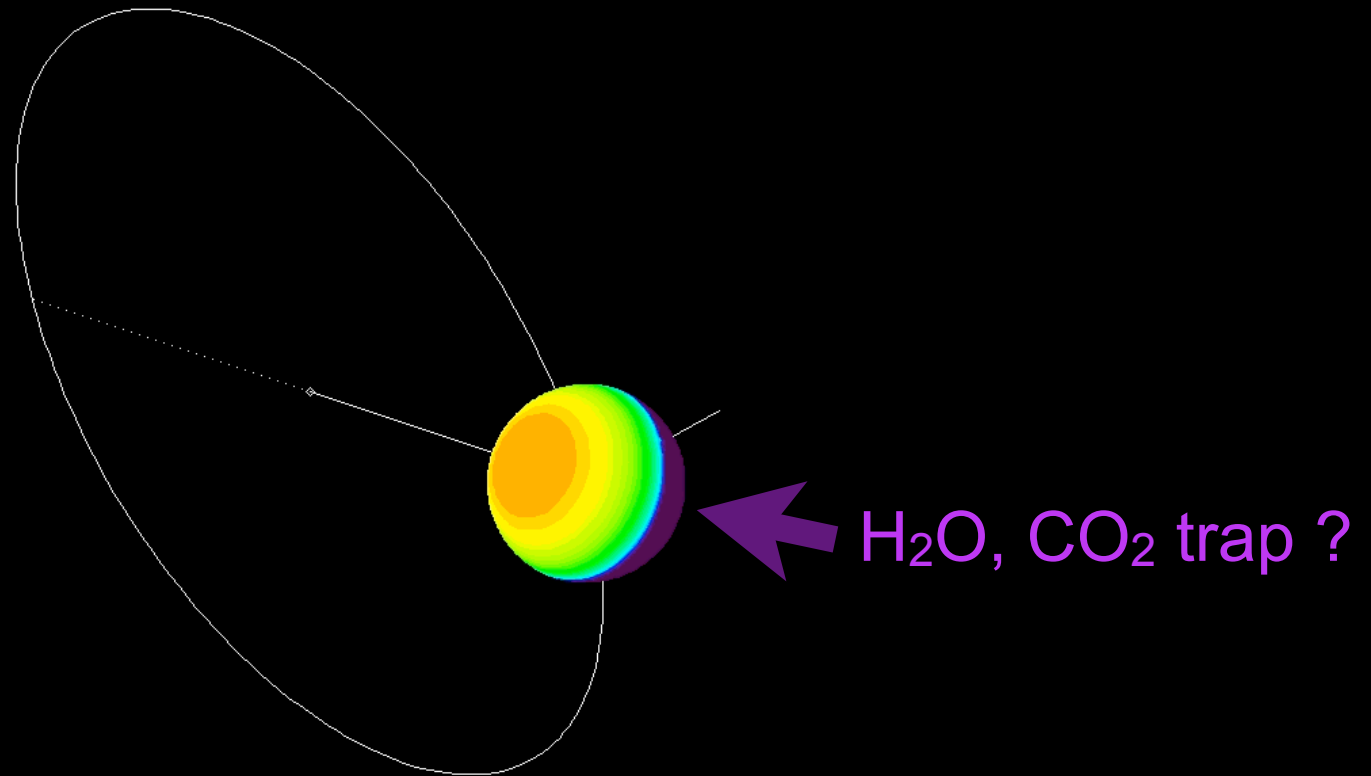
slowly-rotating or synchronous planets

Tidally-evolved planets should be found either in their equilibrium rotation state ($f(e)$) or in a spin-orbit resonance (which could be 1:1)

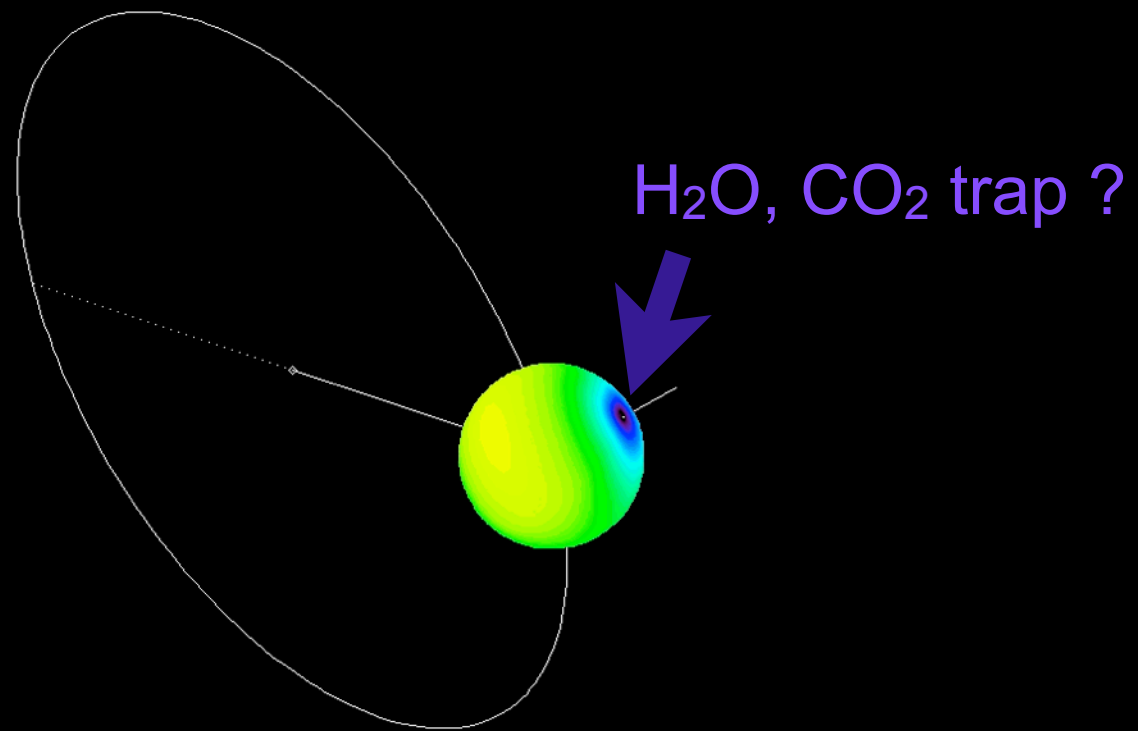
For GJ581d: $P_{\text{ROT}}=0.4P_{\text{ORB}}$ if $e=0.38$

$P_{\text{ROT}}=0.55P_{\text{ORB}}$ if $e=0.3$

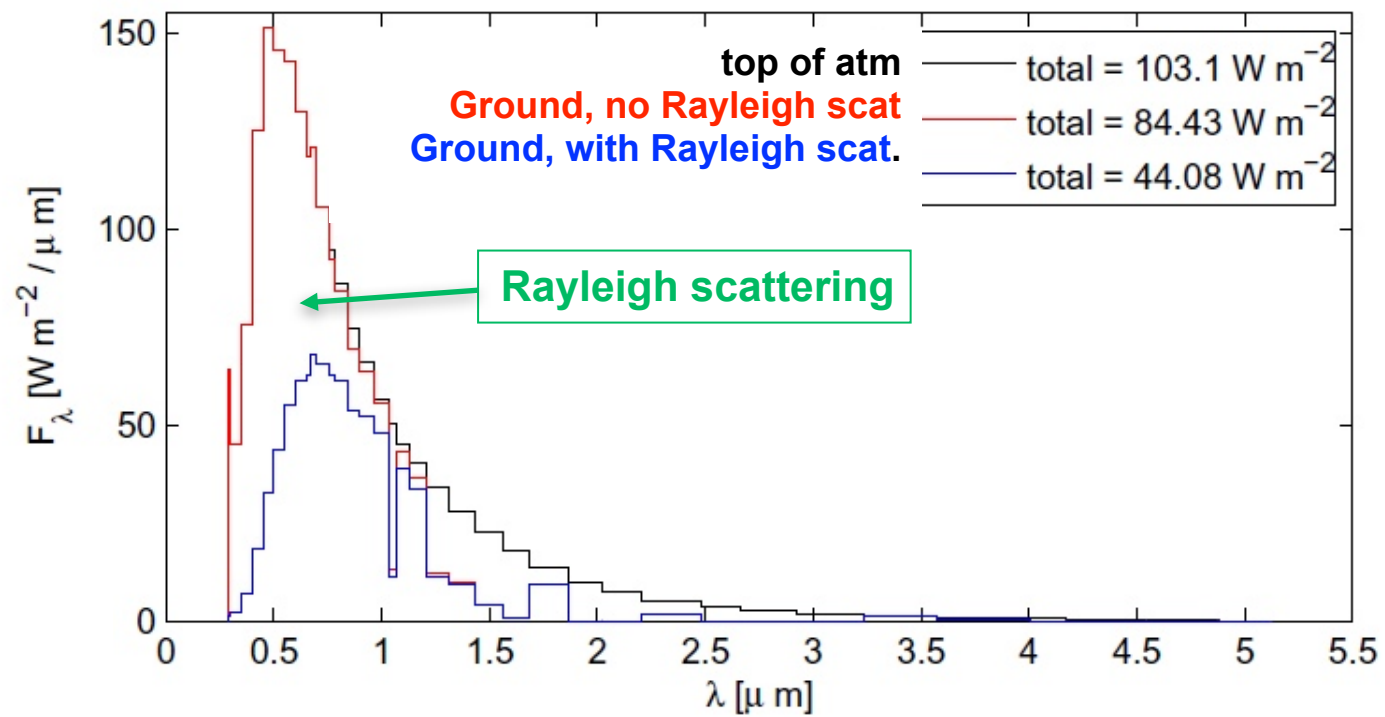
Leconte et al., 2010, Heller et al., 2011



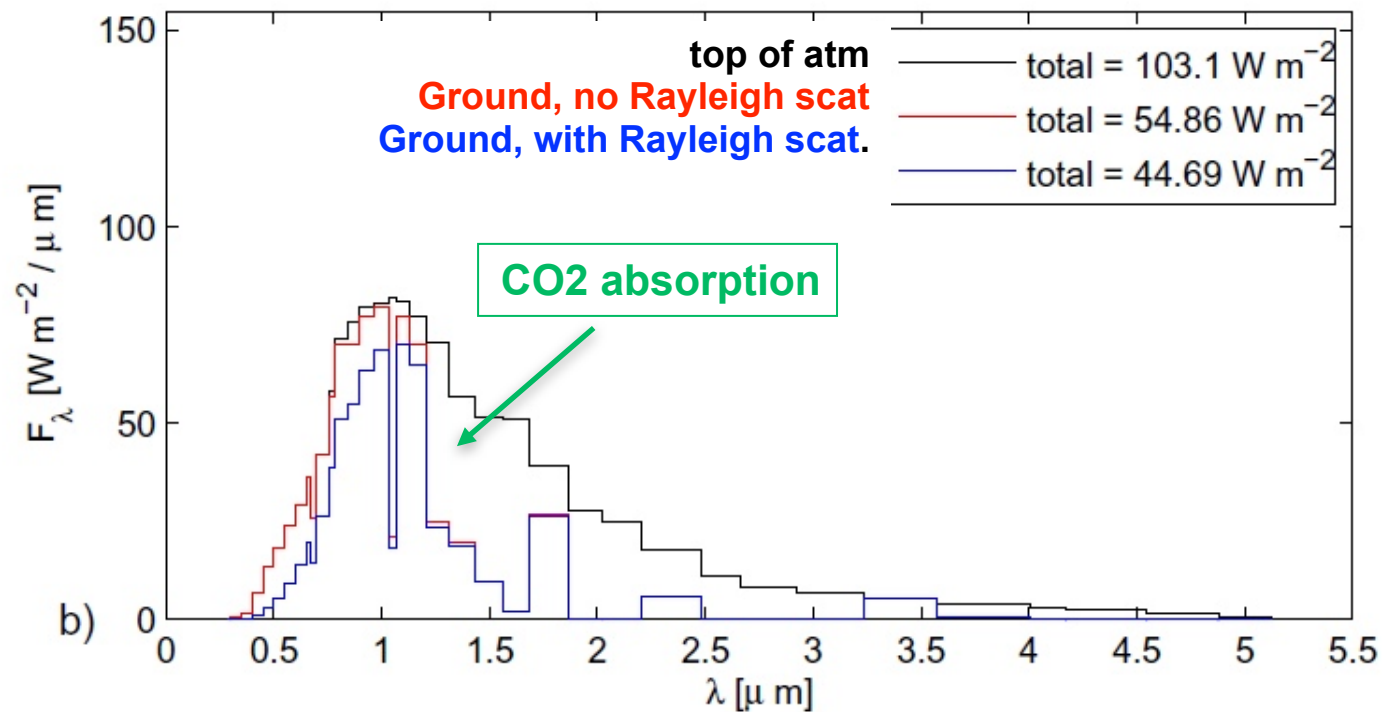
0° obliquity



G star

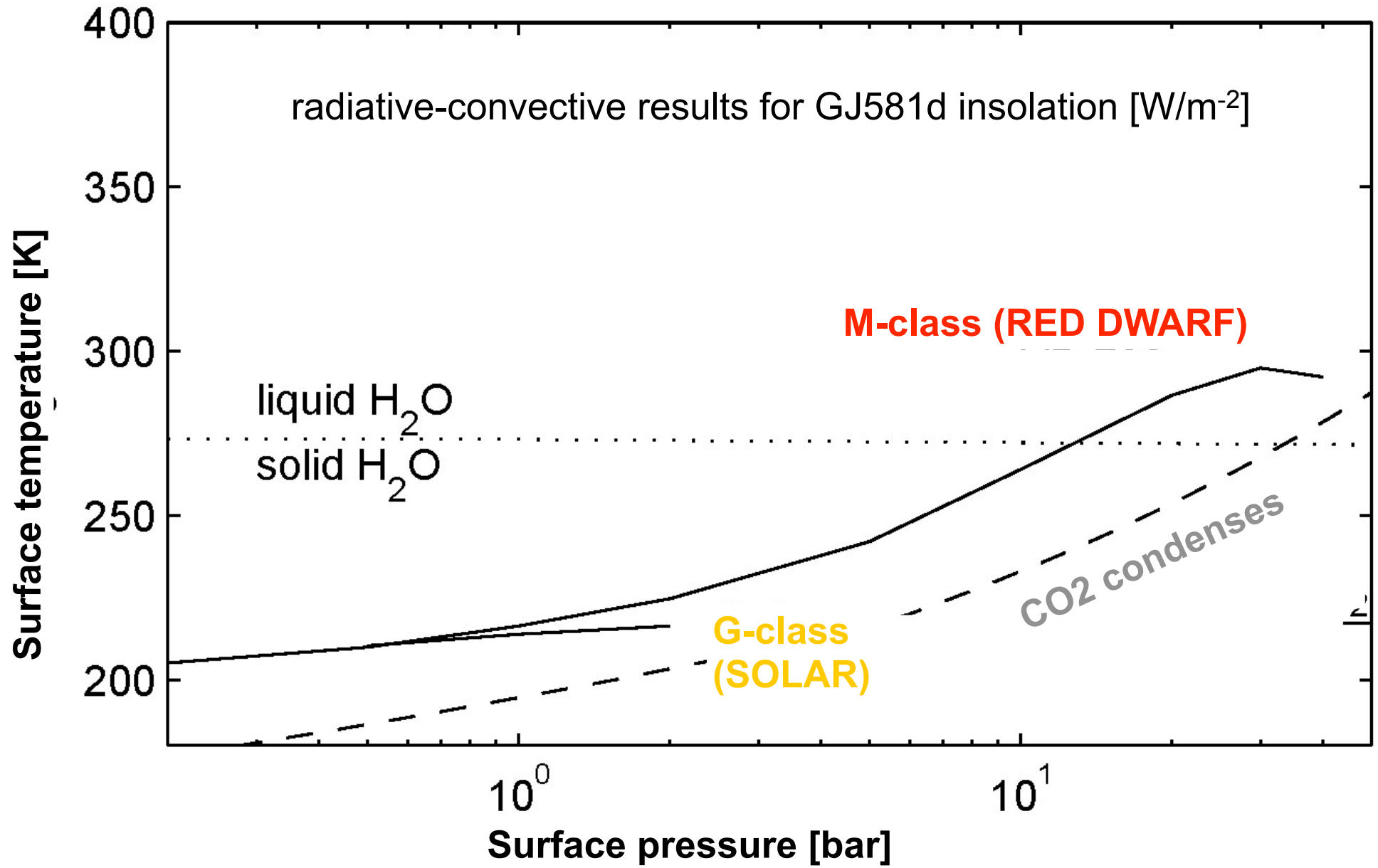


CO₂ atmosphere



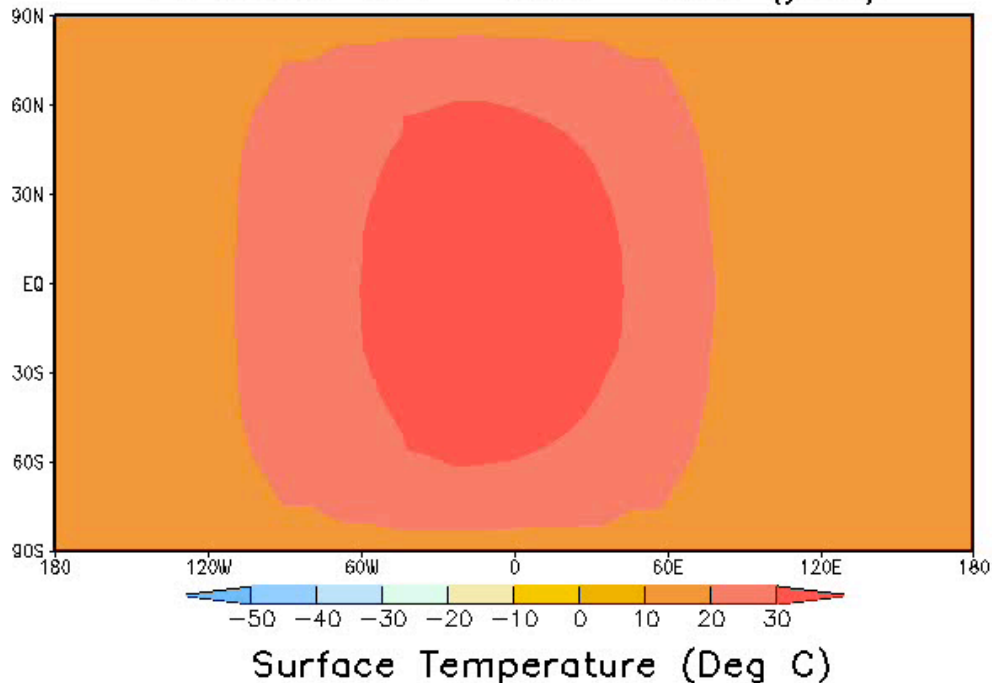
M star

Pure CO₂ atmospheres (no CO₂ clouds)

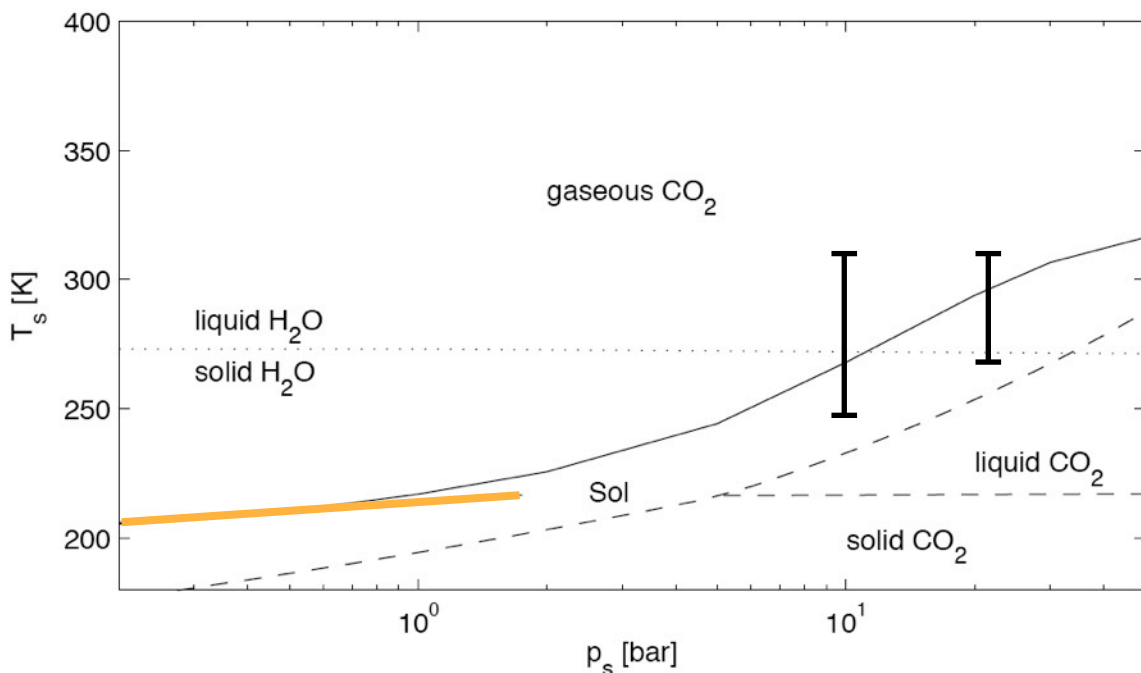
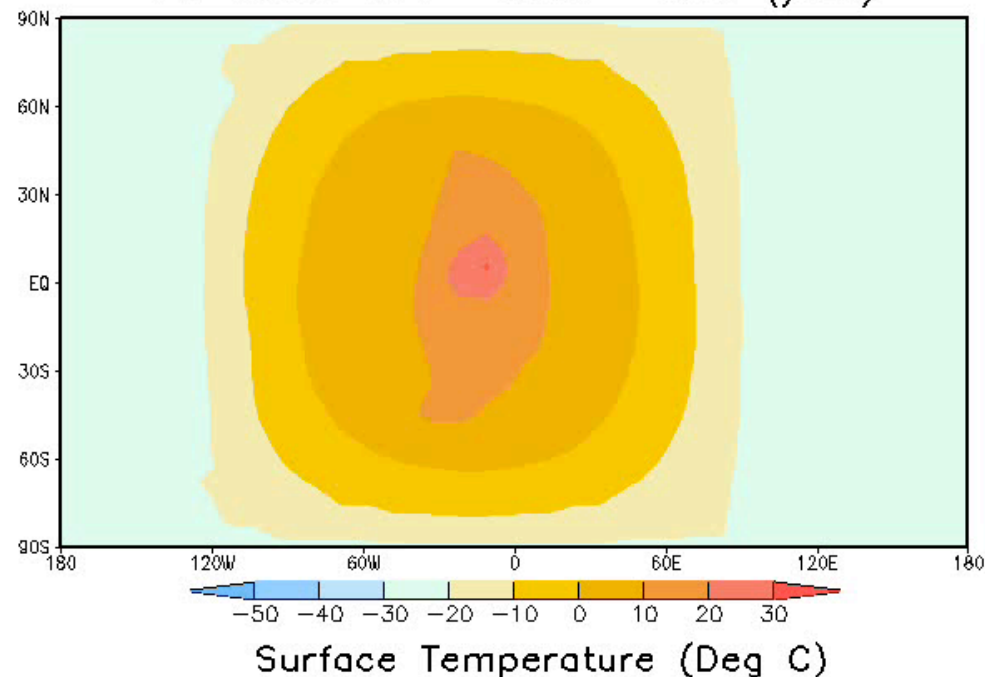


pure CO₂

Ps=20bar n11 Date = 0.00 (year)



Ps=10bar n11 Date = 0.00 (year)

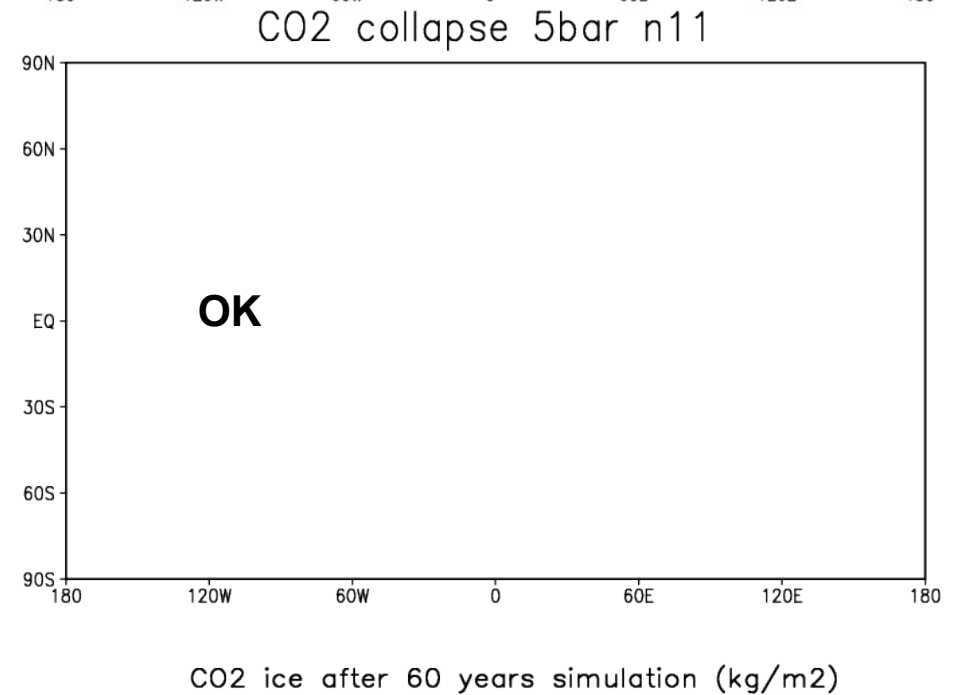
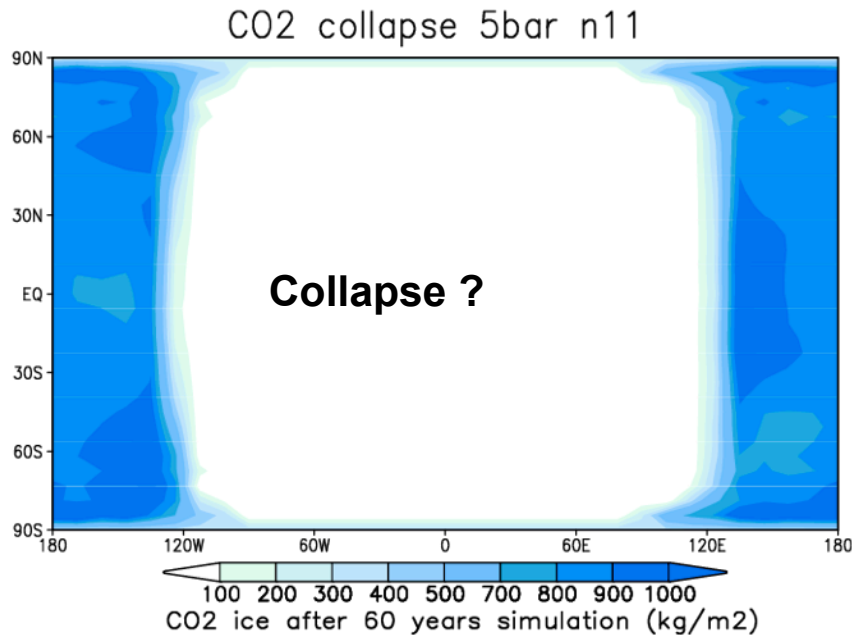
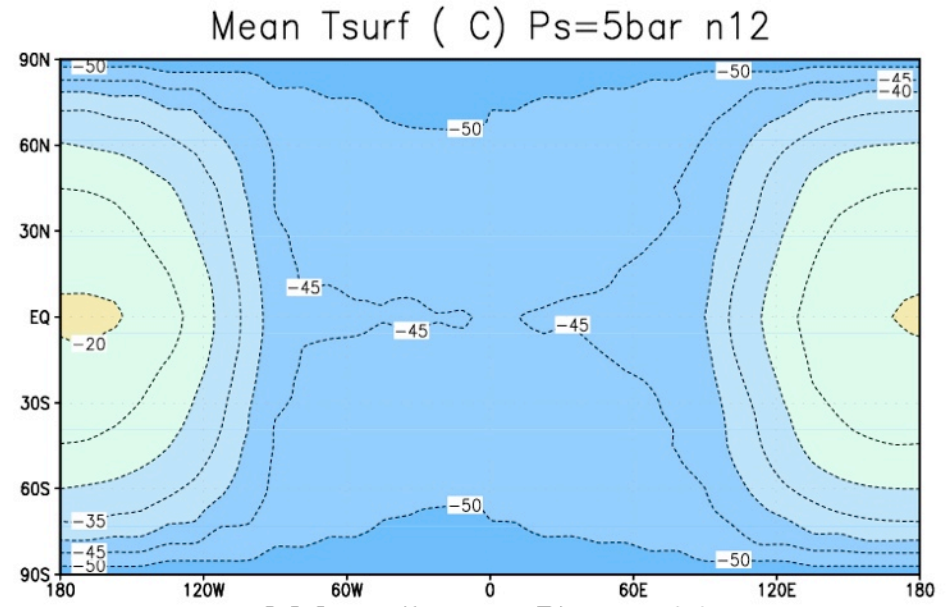
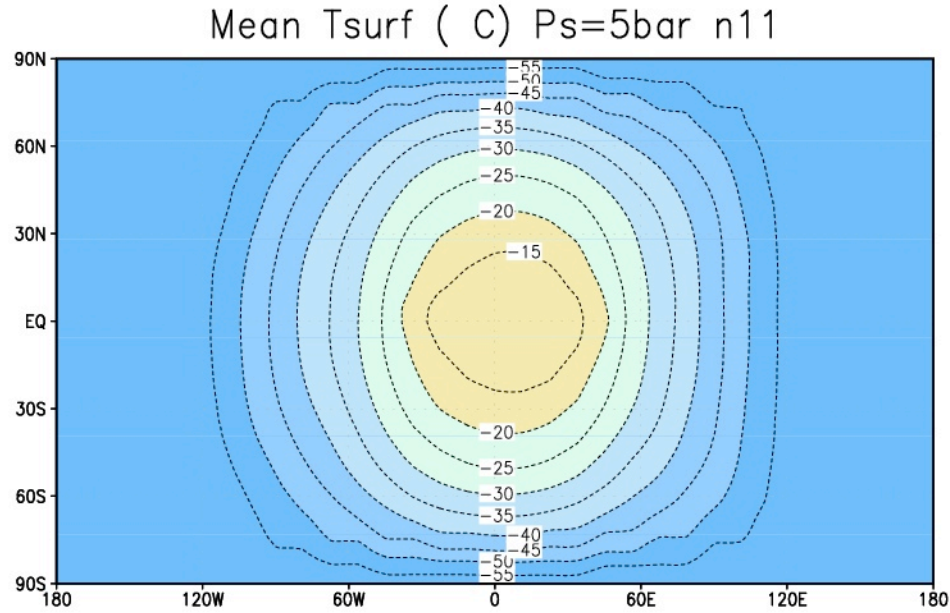


No CO₂ collapse but cold trap for H₂O in the 10 bar case

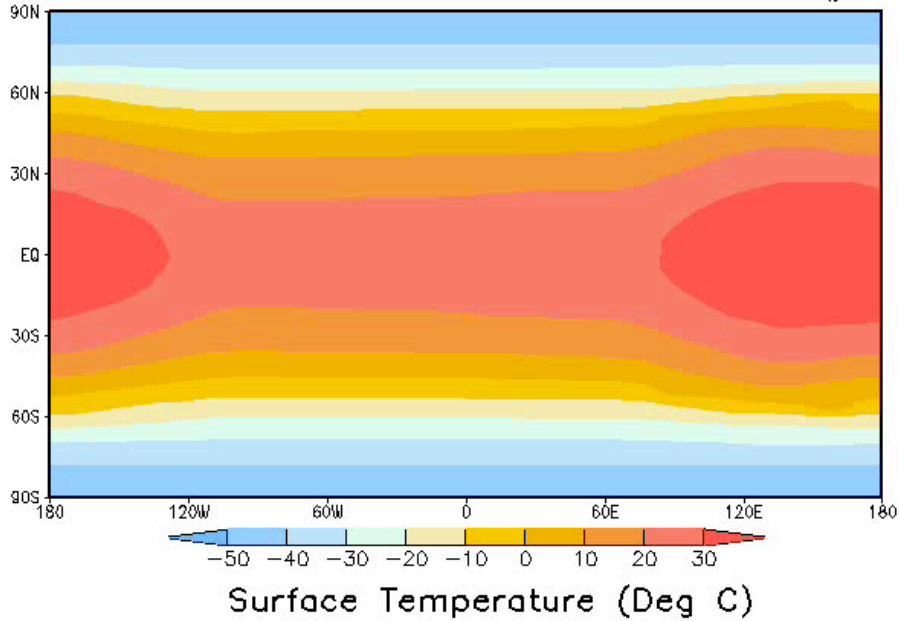
Wordsworth et al., 2011

mercredi 8 juin 2011

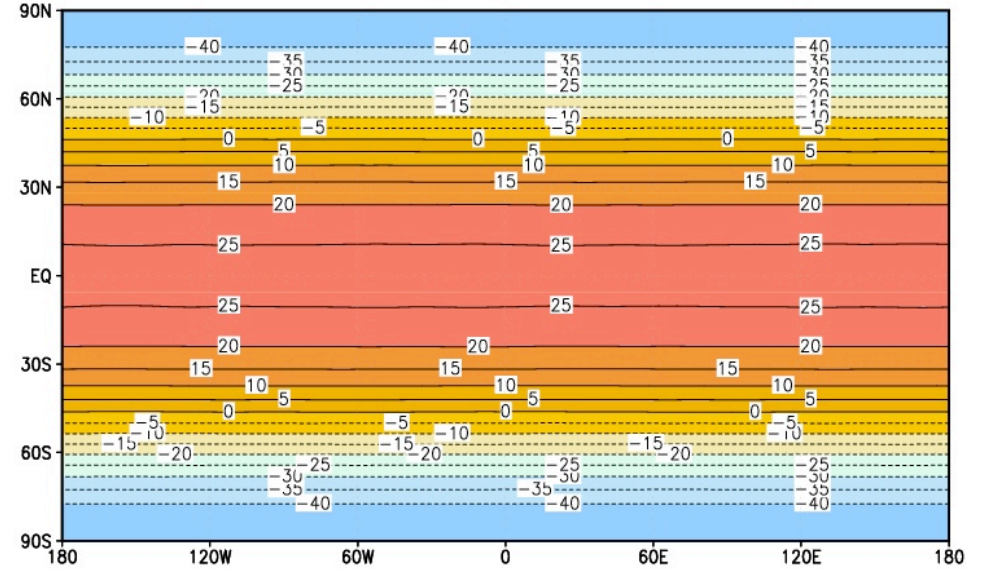
Annual mean temperatures, rotation and CO₂ collapse 5 bars



Ps=10bar Earth-like rotation Date = 0.00 (year)

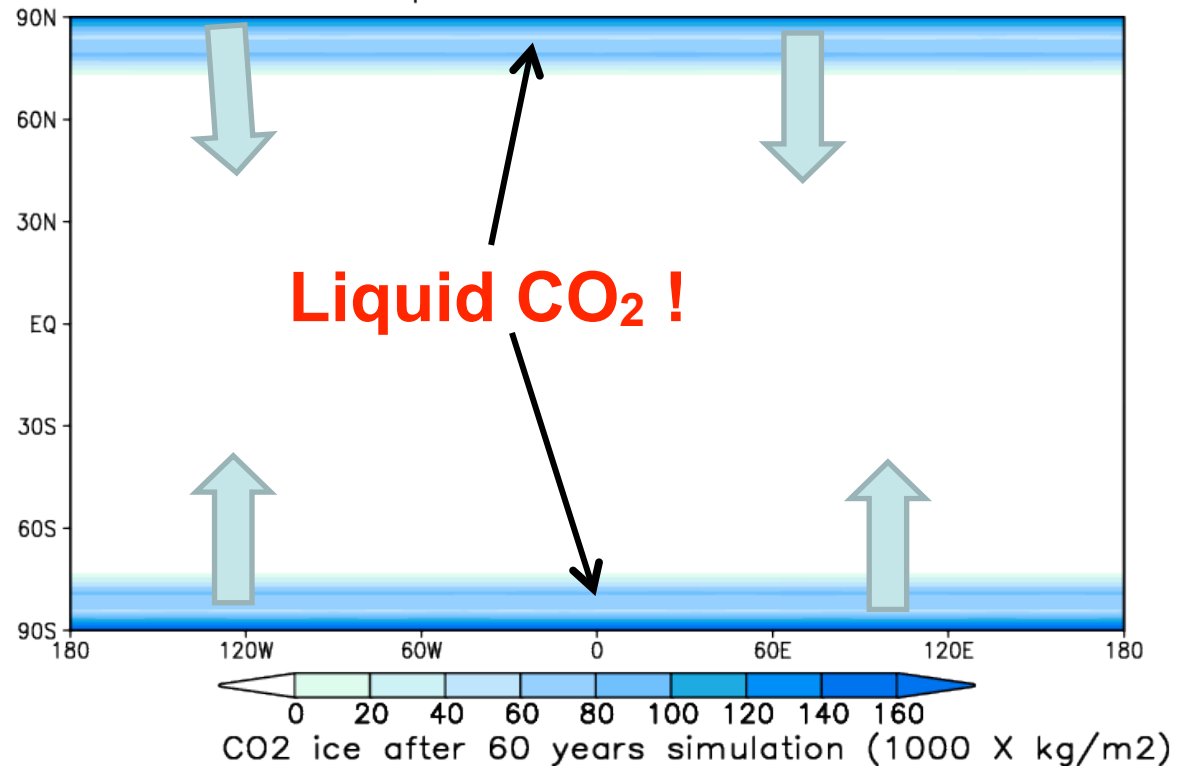


Mean Tsurf (C) Ps=10bar Erot



annual mean temperatures,
rotation and CO₂ collapse
10 bars

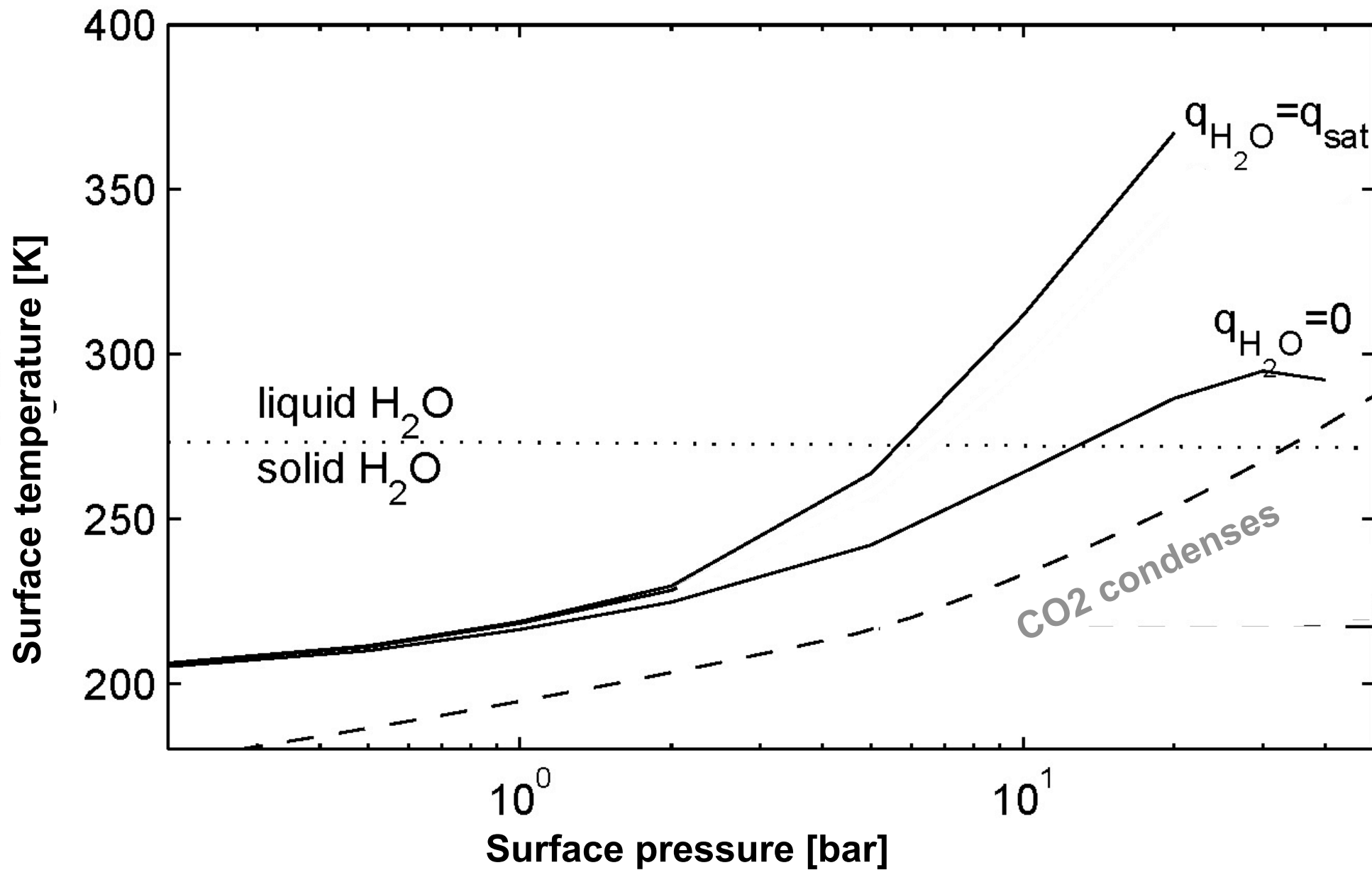
CO₂ collapse 10bar Earth rotation



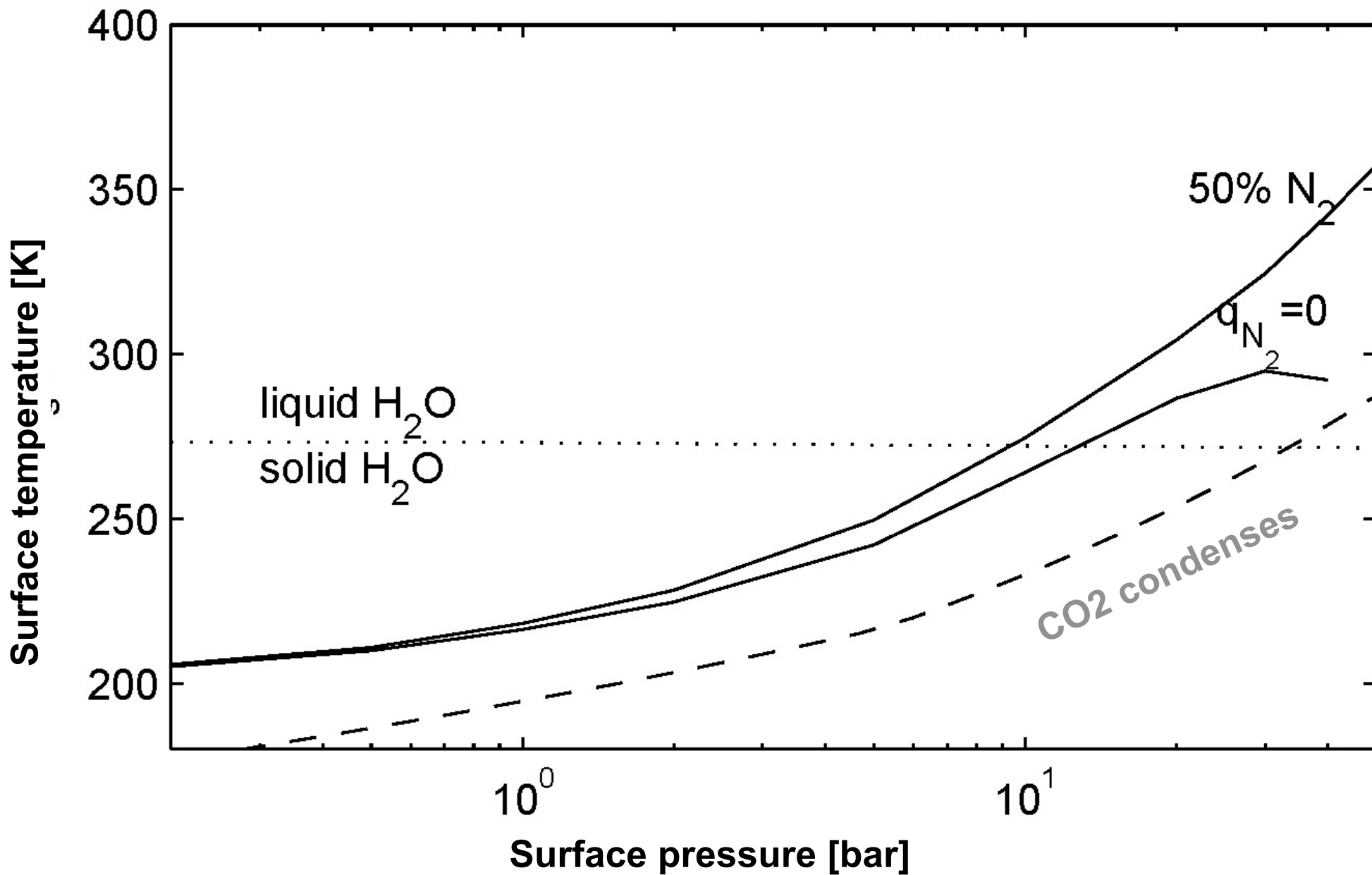
Wordsworth et al., 2011

mercredi 8 juin 2011

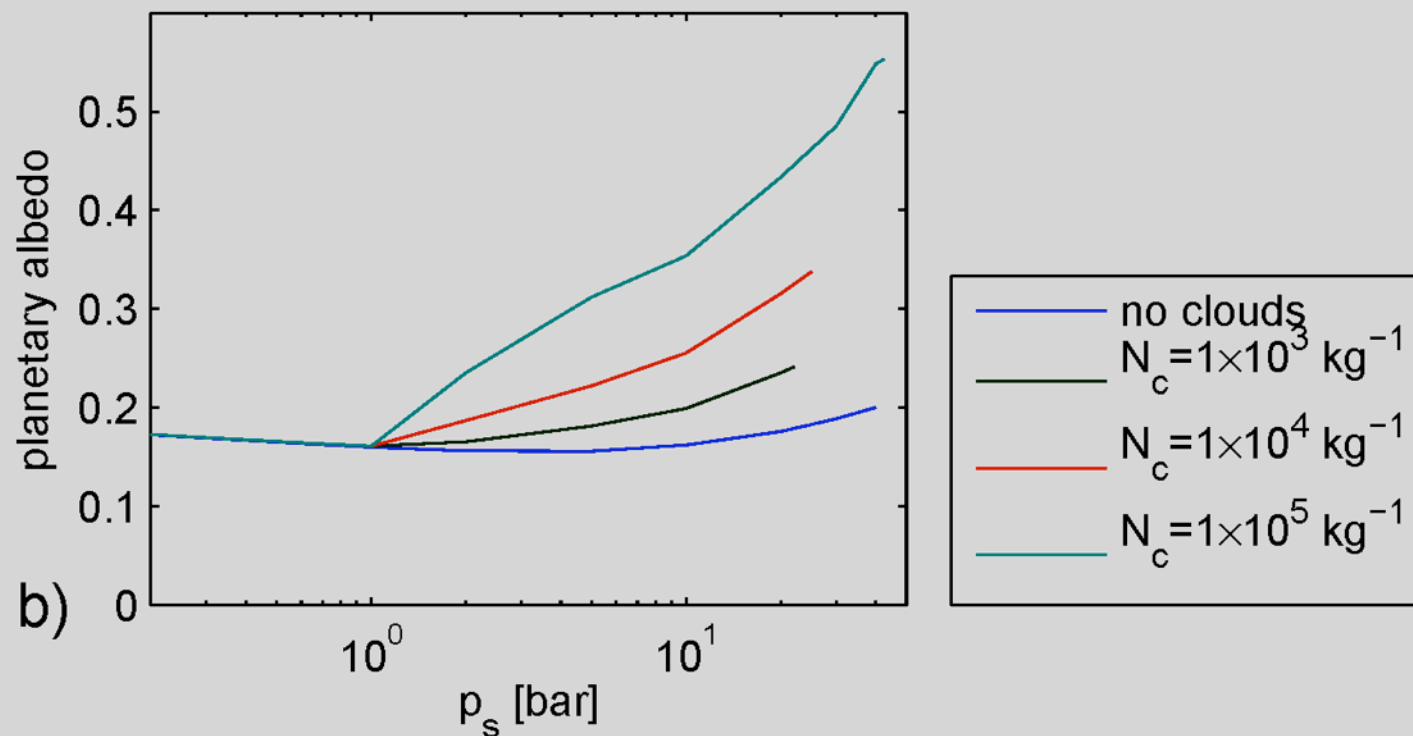
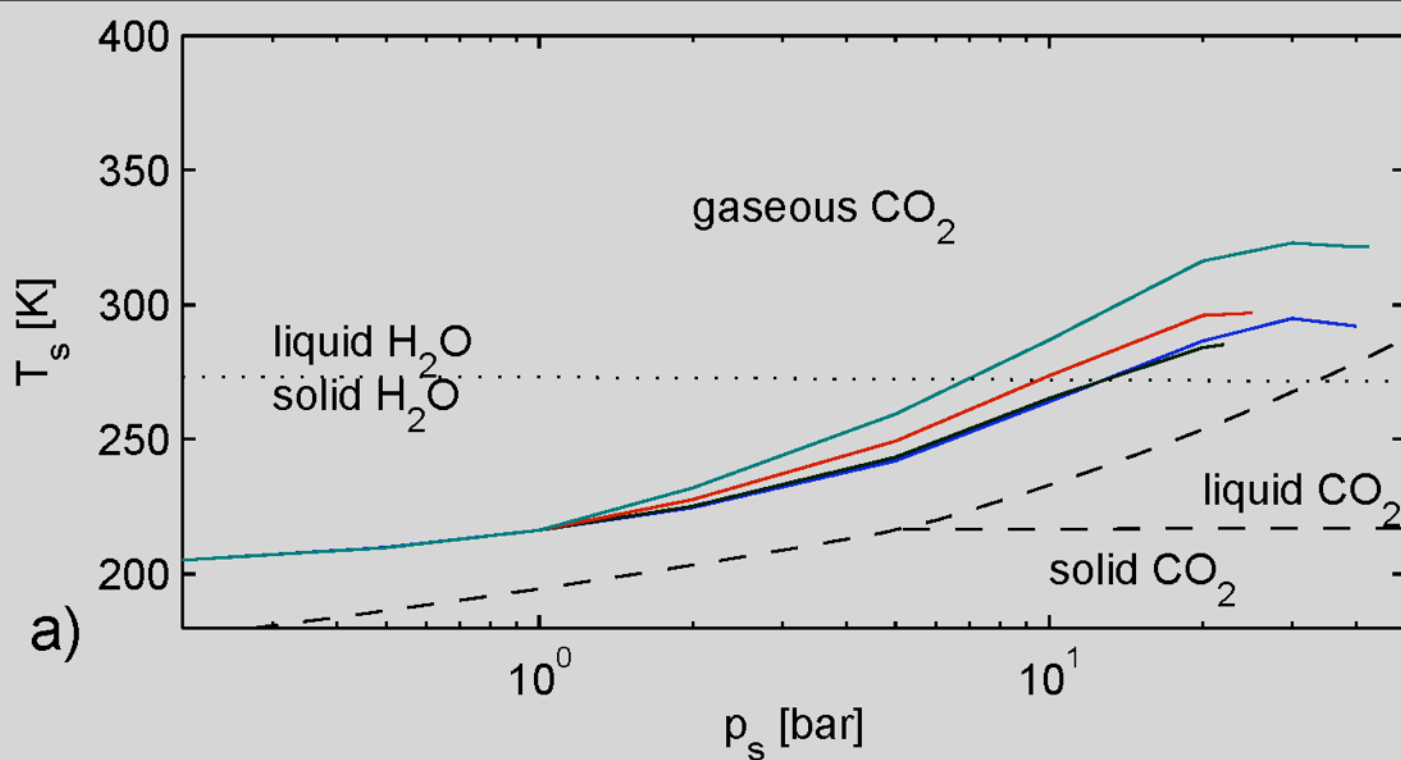
CO₂ + H₂O



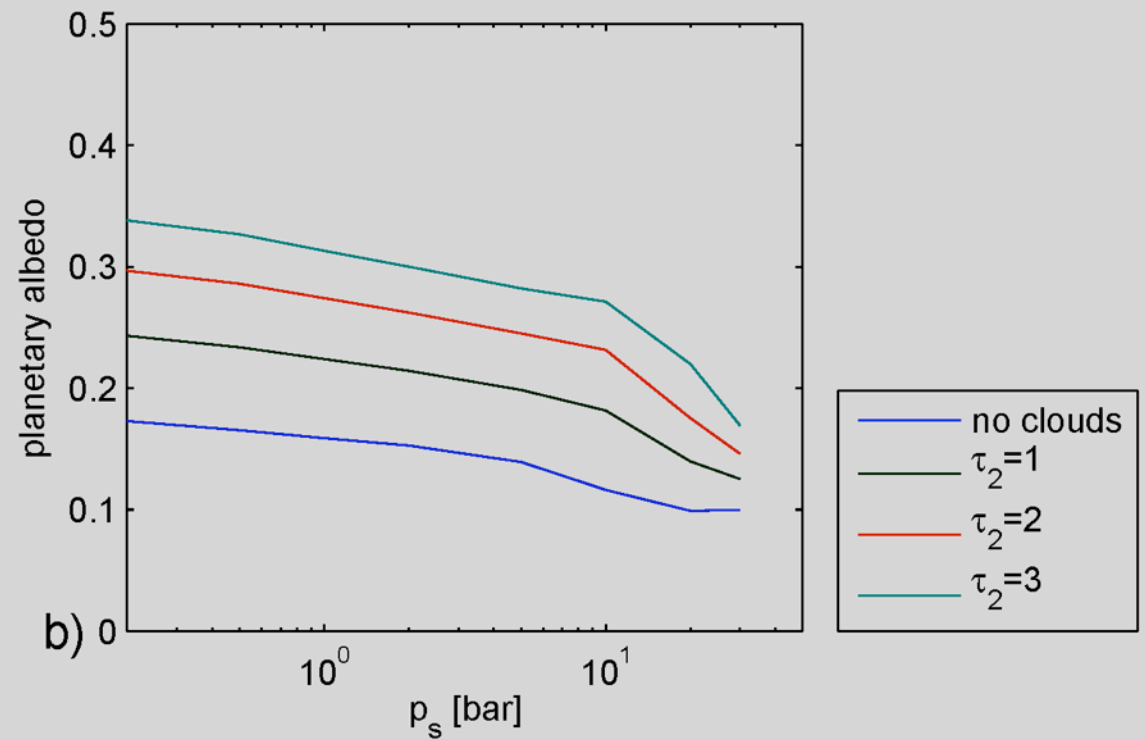
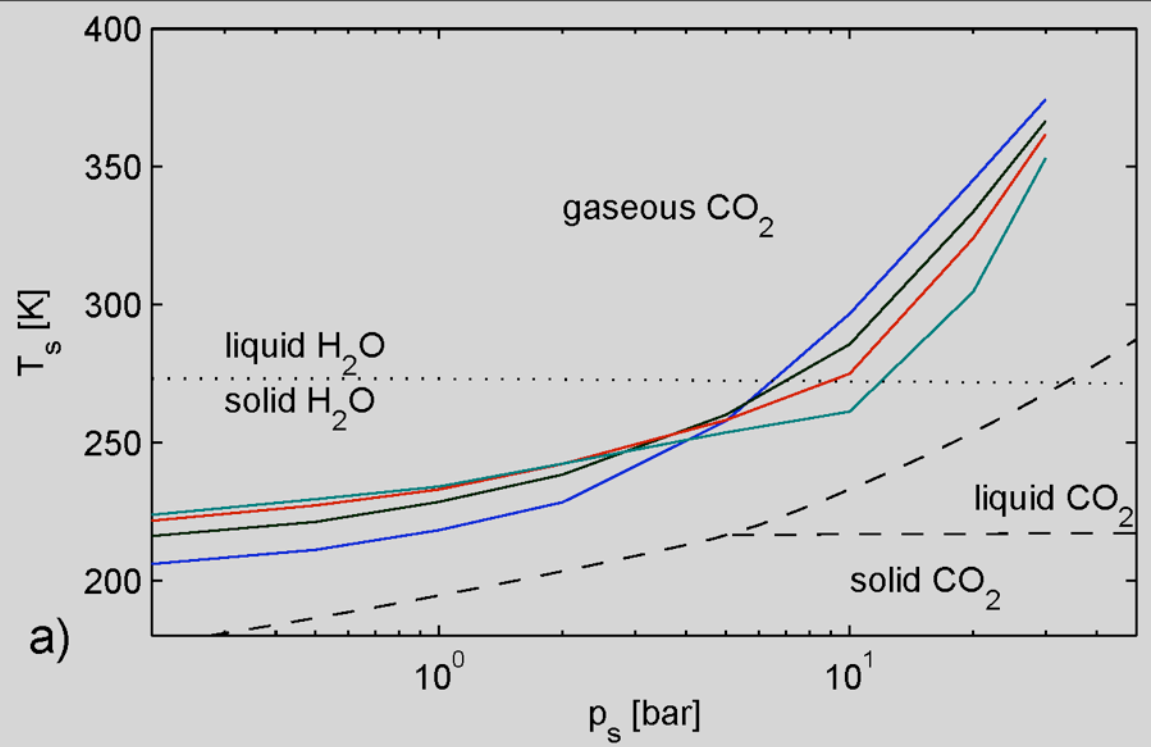
CO₂ + N₂

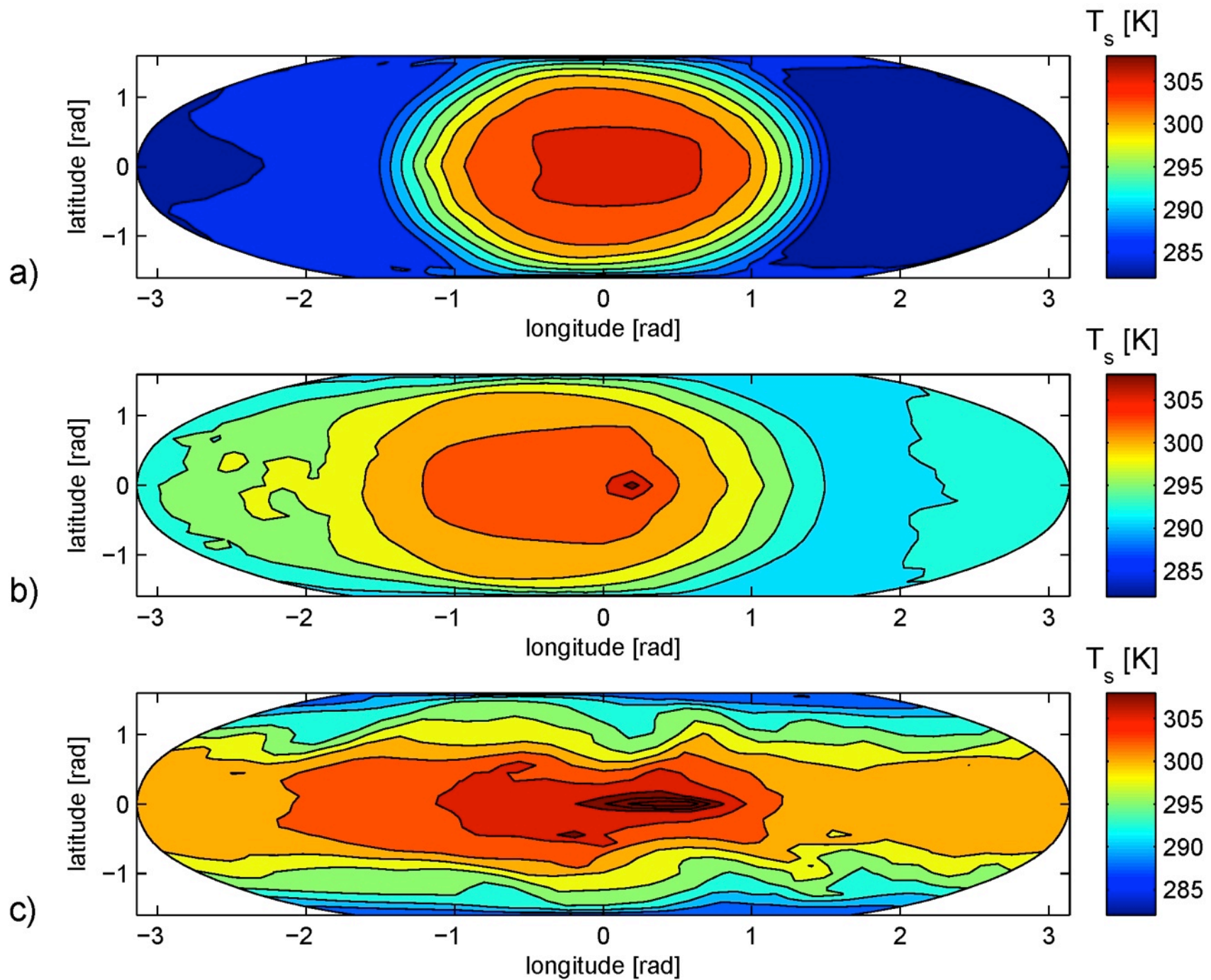


CO₂ clouds



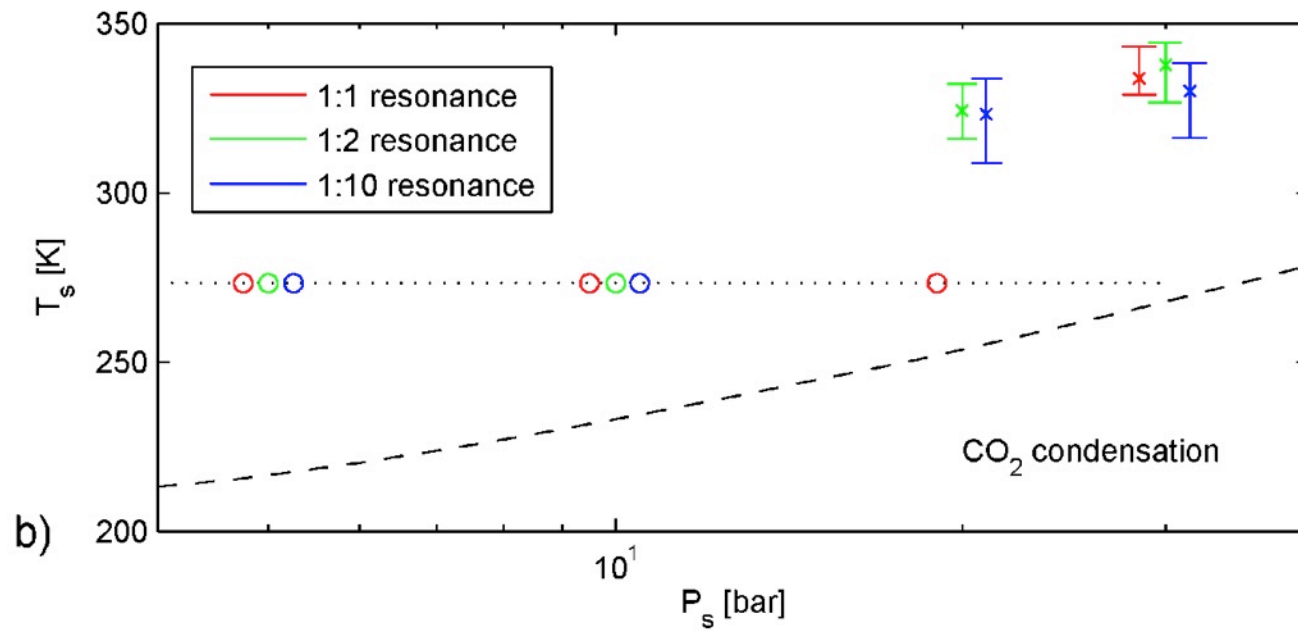
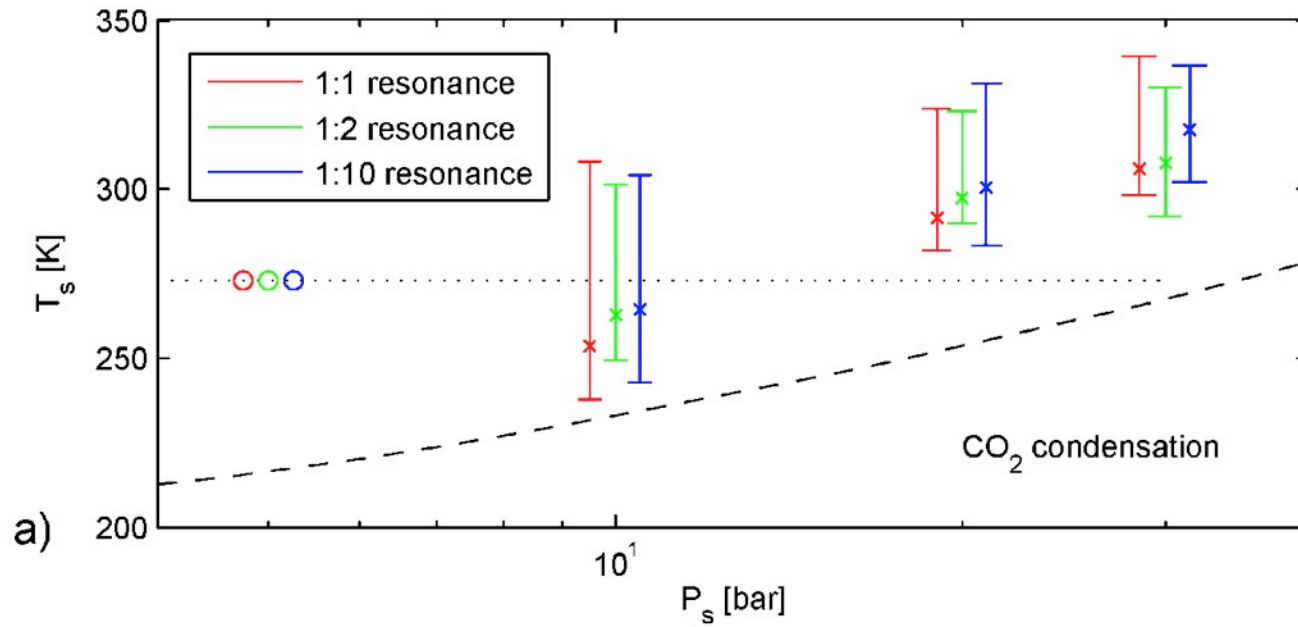
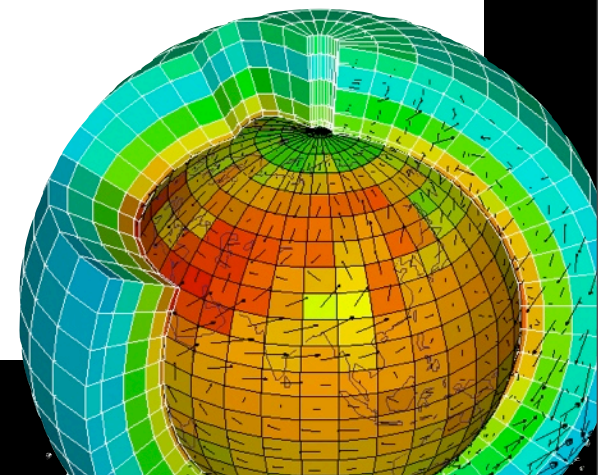
H₂O





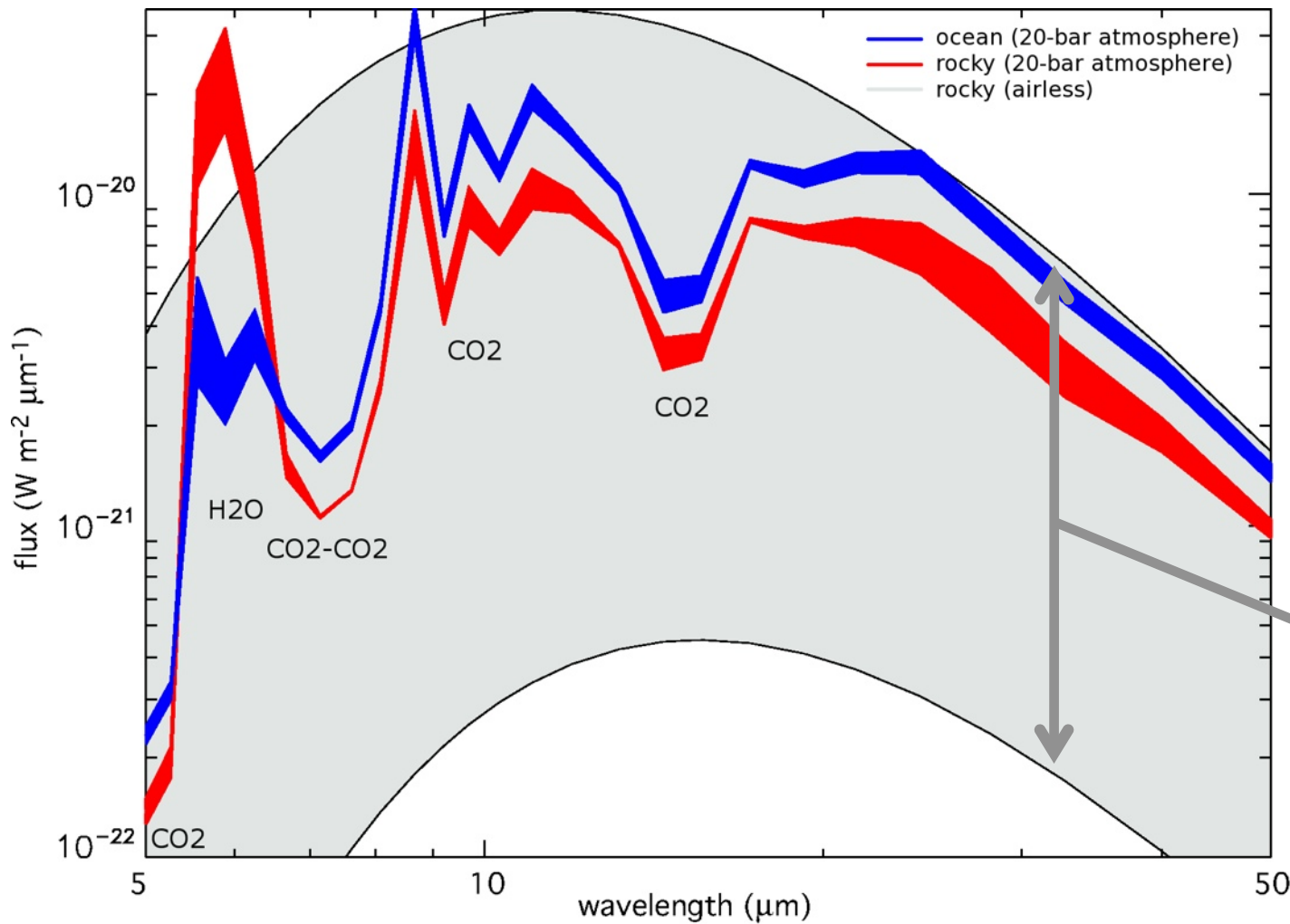
ROCKY

OCEAN



Wordsworth et al. (2011)

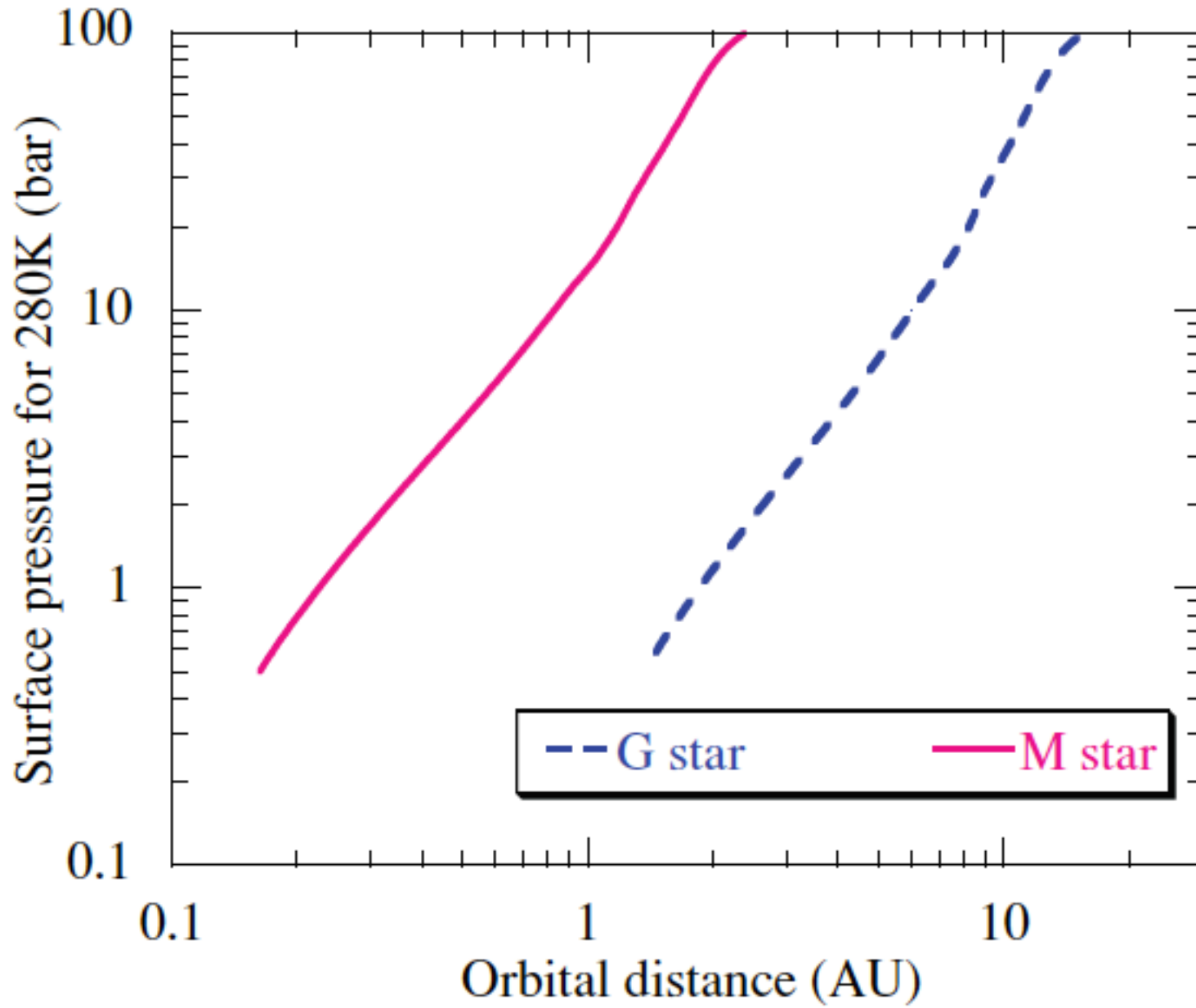
How can we distinguish the various possible cases?



Planet / star
contrast ratio of
order 10^{-6}
→ TPF / Darwin
mission required

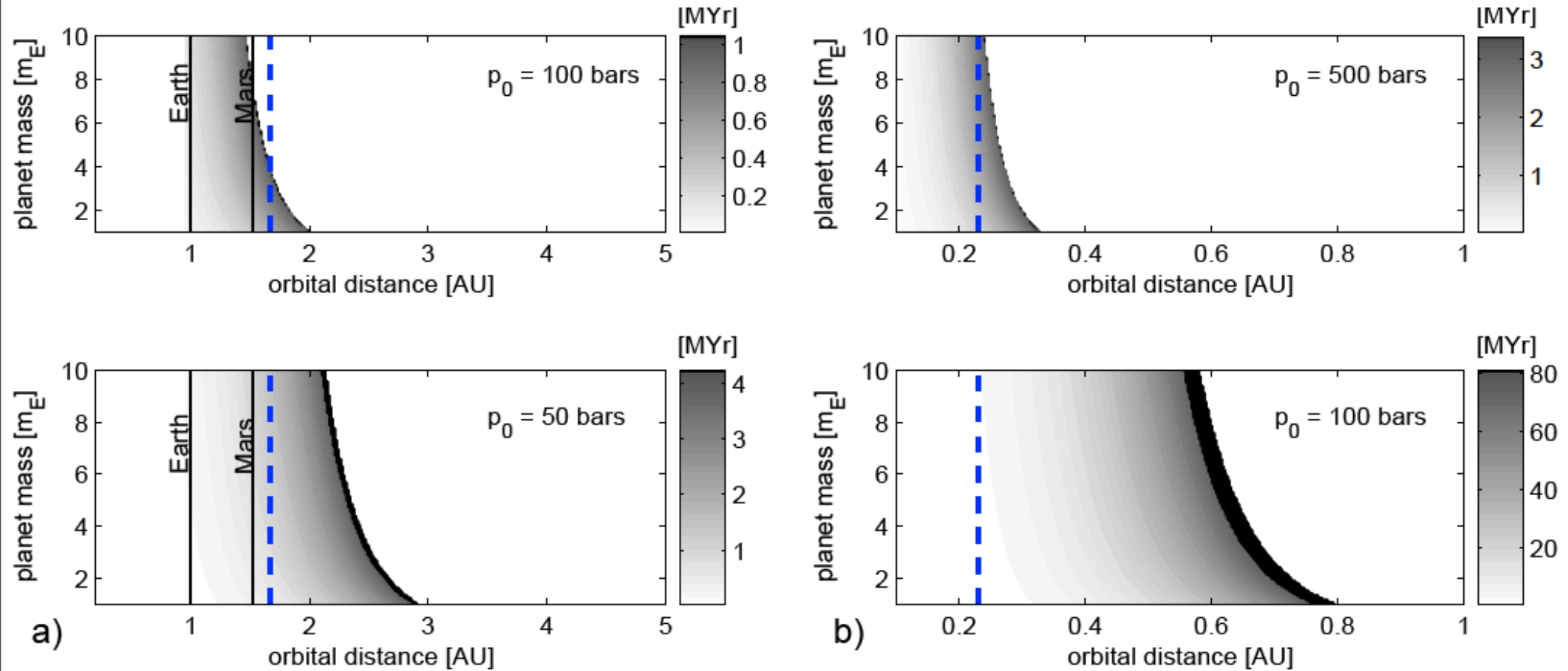
Large orbital
flux variations
in airless case

Greenhouse warming by H₂ atmospheres

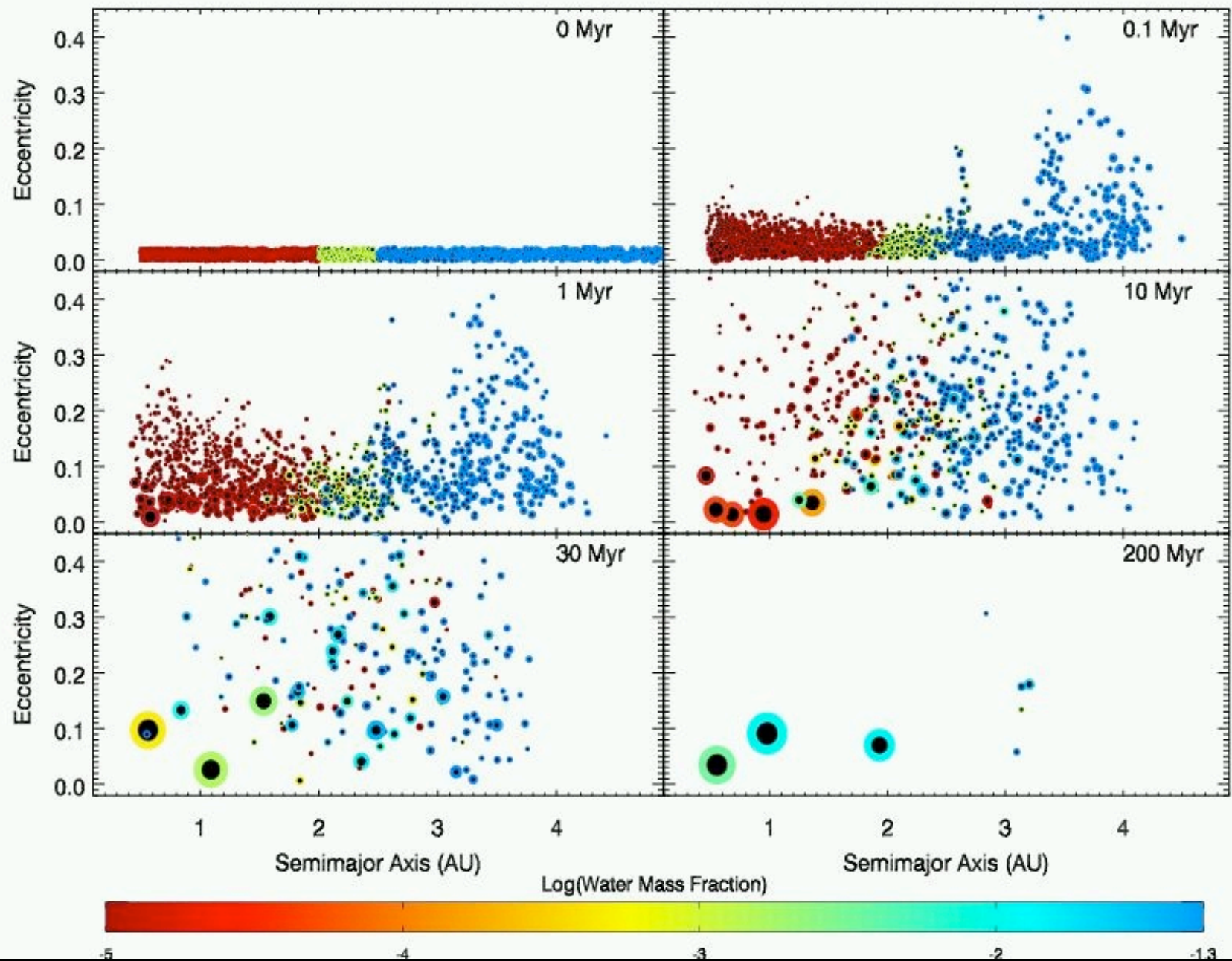


Pierrehumbert and Gaidos, 2011

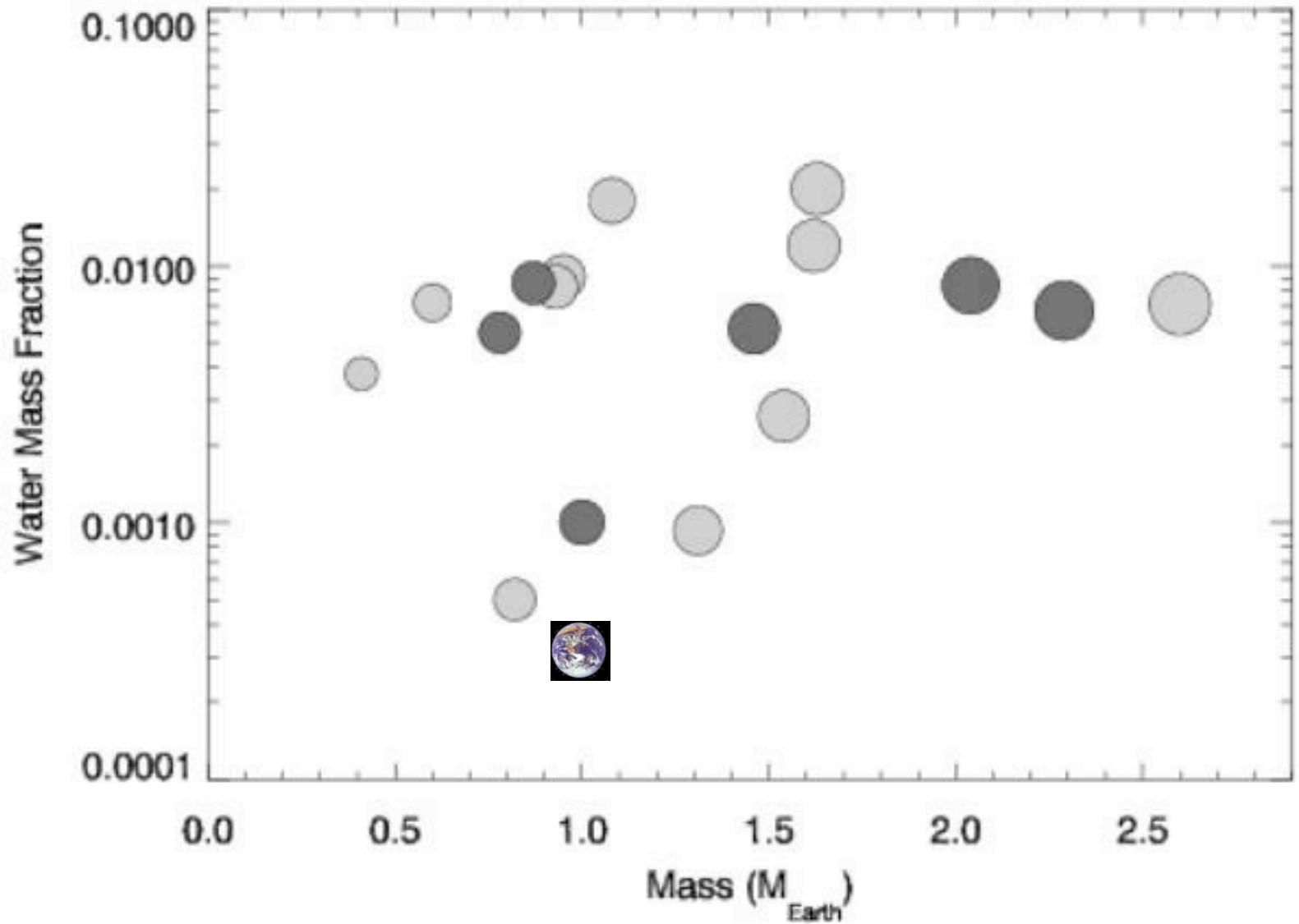
How long does the right H₂ pressure can be sustained ?



Wordsworth, submitted



Raymond et al., 2006



Raymond et al., 2007

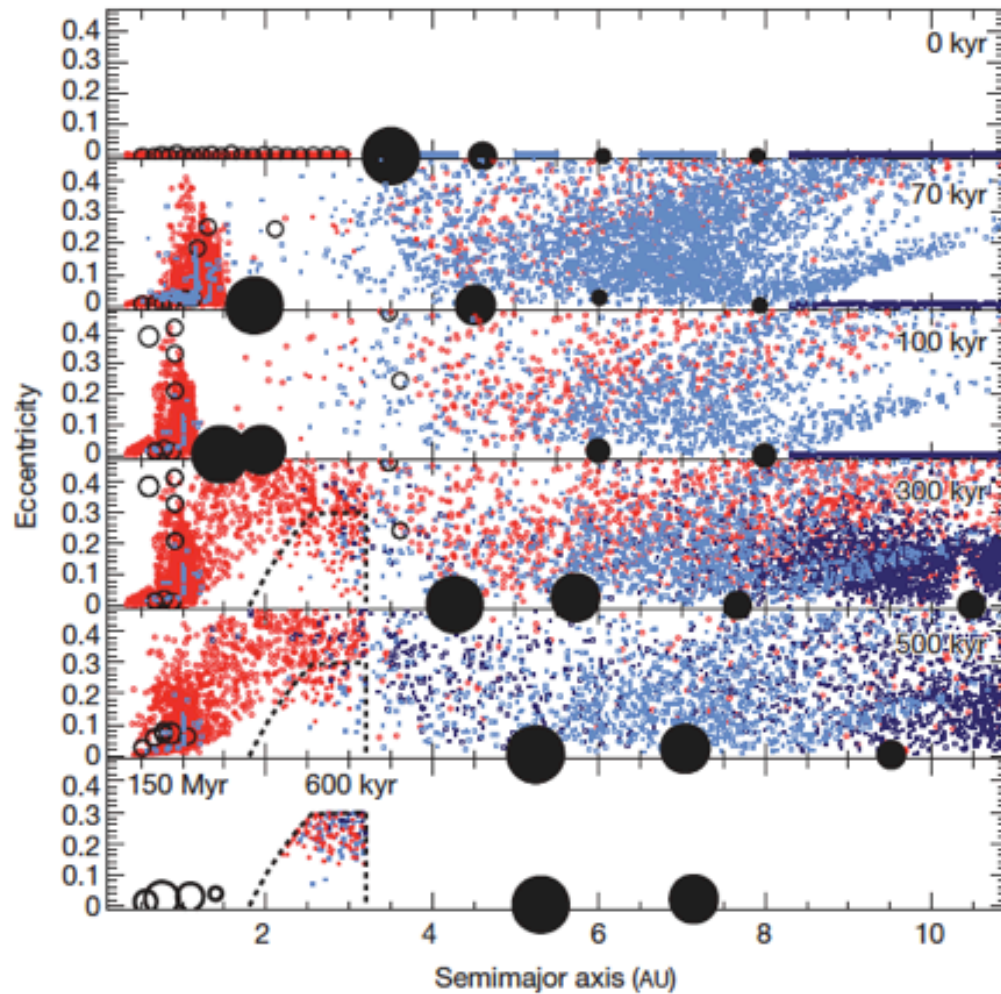


Figure 2 | The evolution of the small-body populations during the growth and migration of the giant planets, as described in Fig. 1. Jupiter, Saturn, Uranus and Neptune are represented by large black filled circles with evident inward-then-outward migration, and evident growth of Saturn, Uranus and Neptune. S-type planetesimals are represented by red dots, initially located between 0.3 and 3.0 AU. Planetary embryos are represented by large open circles scaled by $M^{1/3}$ (but not in scale relative to the giant planets), where M is mass. The C-type planetesimals starting between the giant planets are shown as light blue dots, and the outer-disk planetesimals as dark blue dots, initially between 8.0 and 13.0 AU. For all planetesimals, filled dots are used if they are inside the

and the C types from between the giant planets and from 8.0 to 13.0 AU. The present-day asteroid belt consists of more than just S- and C-type asteroids, but this diversity is expected to result from compositional gradients within each parent population (Supplementary Information). There is a correlation between the initial and final locations of implanted asteroids (Fig. 3a). Thus, S-type objects dominate in the inner belt, while C-type objects dominate in the outer belt (Fig. 3b). Both types of asteroid share similar distributions of eccentricity and inclination (Fig. 3c, d). The present-day asteroid belt is expected to have had its eccentricities and inclinations reshuffled during the so-called late heavy bombardment (LHB)^{13,14}; the final orbital distribution in our simulations matches the conditions required by LHB models.

Given the overall efficiency of implantation of $\sim 0.07\%$, our model yields $\sim 1.3 \times 10^{-3} M_{\oplus}$ of S-type asteroids at the time of the dissipation of the solar nebula. In the subsequent 4.5 Gyr, this population will be depleted by 50–90% during the LHB event^{13,14} and by a further factor of $\sim 2-3$ by chaotic diffusion¹⁵. The present-day asteroid belt is estimated to have a mass of $6 \times 10^{-4} M_{\oplus}$, of which 1/4 is S-type and 3/4 is C-type¹². Thus our result is consistent within a factor of a few with the S-type portion of the asteroid belt.

The C-type share of the asteroid belt is determined by the total mass of planetesimals between the giant planets and between 8 and 13 AU, which are not known a priori. Requiring that the mass of implanted C-type material be three times that of the S-type, and given the implantation efficiencies reported above, this implies that the following

